

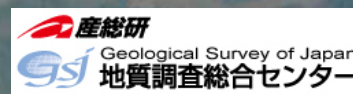
# Hundred-year advances in understanding and surveying volcanic degassing

Patrick Allard<sup>1</sup>, Marie Edmonds<sup>2</sup> and Hiroshi Shinohara<sup>3</sup>

<sup>1</sup> Institut de Physique du Globe de Paris (IPGP), Paris University, France

<sup>2</sup> Department of Earth Sciences, Cambridge University, UK

<sup>3</sup> Geological Survey of Japan (GSJ), AIST, Tsukuba, Japan



# Gas release: a systematic manifestation of active volcanoes ! (erupting AND dormant ones)

- Dormant (closed-conduit) volcanoes: **H<sub>2</sub>S**-rich LT fumarolic emissions



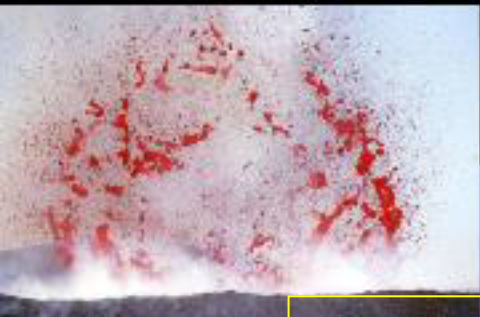
- Open-conduit volcanoes: persistent **SO<sub>2</sub>**-prevalent HT degassing (plumes)



- Erupting volcanoes: discrete **SO<sub>2</sub>**-rich gas emissions



# Motivation



$SO_2$ ,  $HF$ ,  $CO_2$ ,  
 $CO$ ,  $H_2S$ ,  $HCl$ ,  
 $BrO$ , acid aerosol,  
trace metals...

# Global geochemical budgets

Origins and fate (cycles)  
of volatiles on Earth

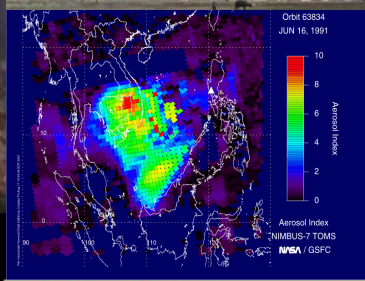
A number of good motivations to study and monitor volcanic gas emissions

# Forecasting

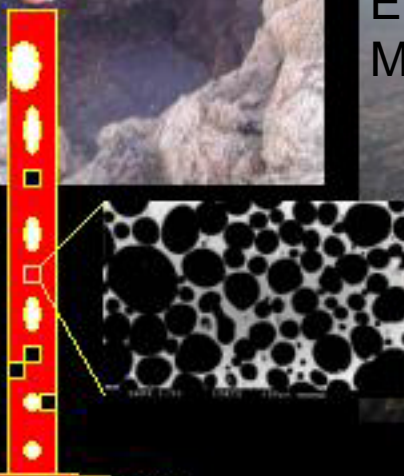
Eruption precursors  
Mitigating hazards



Impacts on the climate



# Magma dynamics



# Health/environmental impacts

e.g., Kilauea, Miyake-jima

# However, studying volcanic degassing has long received poor attention in early developing volcanology



Two main reasons:

- a) Field and technical challenges  
(*delicate access to degassing sites, risks*)
- b) Information relevance ?

Magmatic volatiles considered negligible components by most igneous petrologists since making only a few weight % of magmas



## Magmatic volatiles: « the Maxwell's devils » (N.L. Bowen, 1928)

Only a few rare scientists early intuited a key role of magmatic volatiles in volcanic processes  
(*Albert Brun, Reginald Daly, Frank Perret, Alfred Lacroix*)

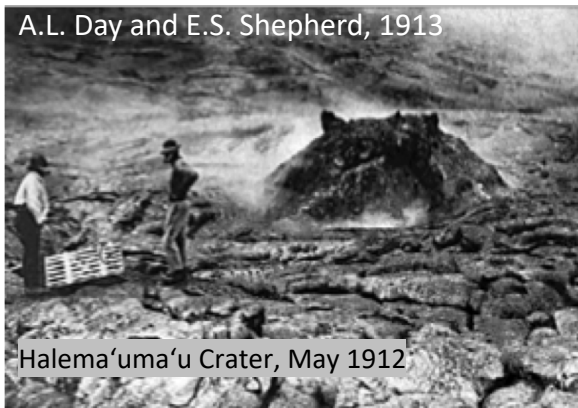


Now, that a heavy liquid should be mobile need not cause surprise...this is to be found, I believe, in its high gas content.  
– *Frank Perret (1913), Lava Fountains of Kilauea, Am. J. Sci.*

“Gas is the active agent, and the magma is its vehicle”

(*F.A. Perret, 1924: The 1906 Vesuvius eruption*)

# Field studies of volcanic gases started to gradually develop in 20<sup>th</sup> century thanks to some obstinated pioneers



A.L. Day and E.S. Shepherd, 1913

Halema'uma'u Crater, May 1912

*Discovery of water predominance in volcanic gas (!)*



Thomas A. JAGGAR, USGS. Kilauea, 1940

*"..gas chemistry is the heart of the volcano magma problem."  
(Jaggar, "Magmatic Gases", 1940)*



Ludovico SICARDI, Italy



Vulcano, Italy

*The first measurements of SO<sub>2</sub> and H<sub>2</sub>S ratio in fumarolic discharges (NH<sub>4</sub>OH-AgNO<sub>3</sub> filled bottles; Sicardi, 1955)*



Showa-shinza, Japan

Sadao Matsuo

Masaaki KAMATA



Sadao MATSUO  
Japan



Tolbachik 1975

Igor MENYAILOV, Kamchatka

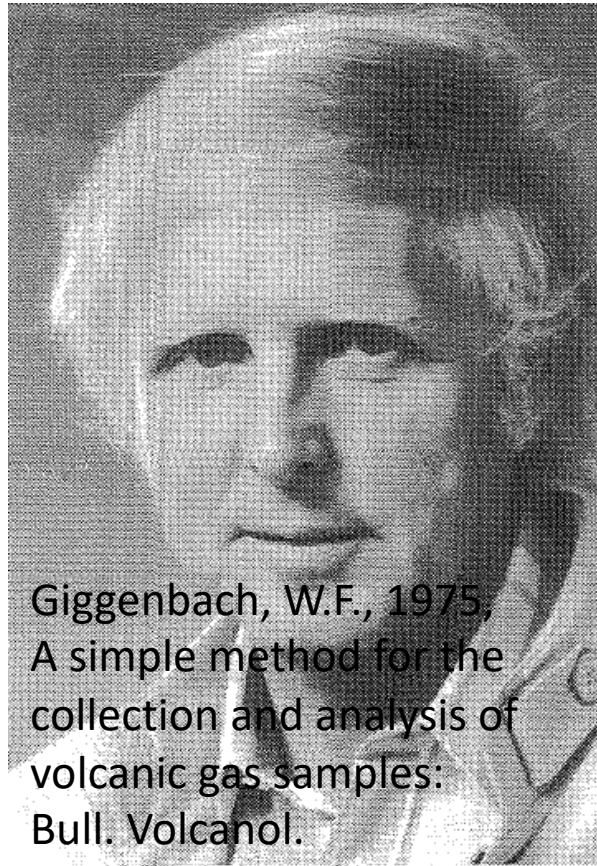
Gudmundur SIGVALDASON, Iceland



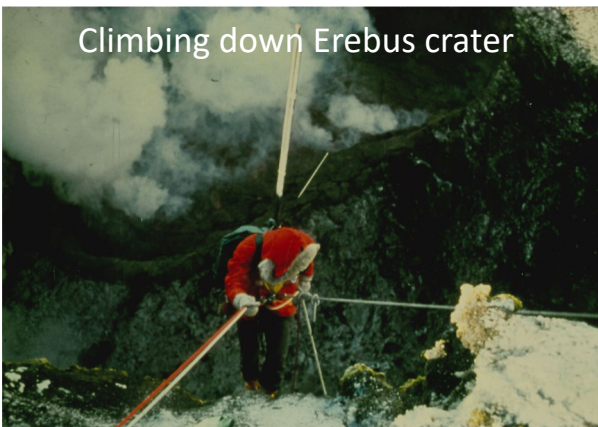
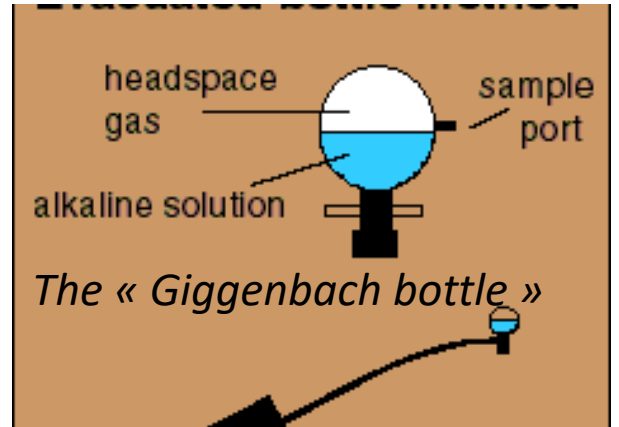
Nyiragongo

Haroun TAZIEFF, France





Giggenbach, W.F., 1975,  
A simple method for the  
collection and analysis of  
volcanic gas samples:  
Bull. Volcanol.



Climbing down Erebus crater

A very special tribute to Werner F Giggenbach  
New Zealand (1937-1997)

*Giggenbach 1996, Chemical composition of volcanic gases*

# Progress from combined studies of volcanic gases (chemistry, thermodynamics, isotopes) and laboratory analysis of volatiles dissolved in magmas

## Gas sampling (chemistry, isotopes)



$H_2O$ ,  $H_2$

$CO_2$ ,  $CO$

$SO_2$ ,  $H_2S$

$HCl$ ,  $HF$ ,  $HBr..$

$N_2$ ,  $He$ ,  $Ar...$

Volatile trace  
metals

( $Po$ ,  $Bi$ ,  $Se$ ,

$Hg$ ,  $Cd$ ,  $As$ ,

$Ag$ ,  $Pb$ ,  $Au$ ,

$Tl$ ,  $Cu$ ,  $Zn...$ )



REVIEWS in  
MINERALOGY  
Volume 30

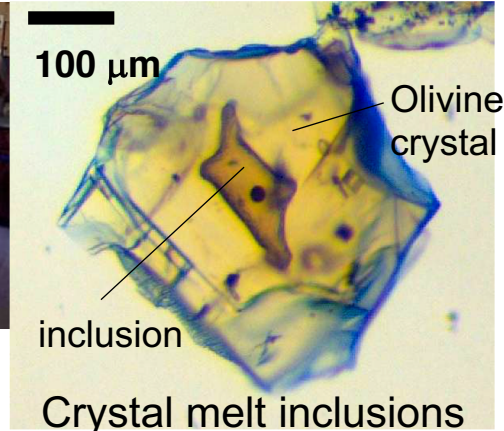
## VOLATILES IN MAGMAS

M.R. Carroll & J.R. Holloway  
*Editors*



Series Editor: Paul H. Ribbe  
MINERALOGICAL SOCIETY OF AMERICA

## Volatiles content & solubility in magmas



## Petrologic experiments volatile solubilities HP-HT

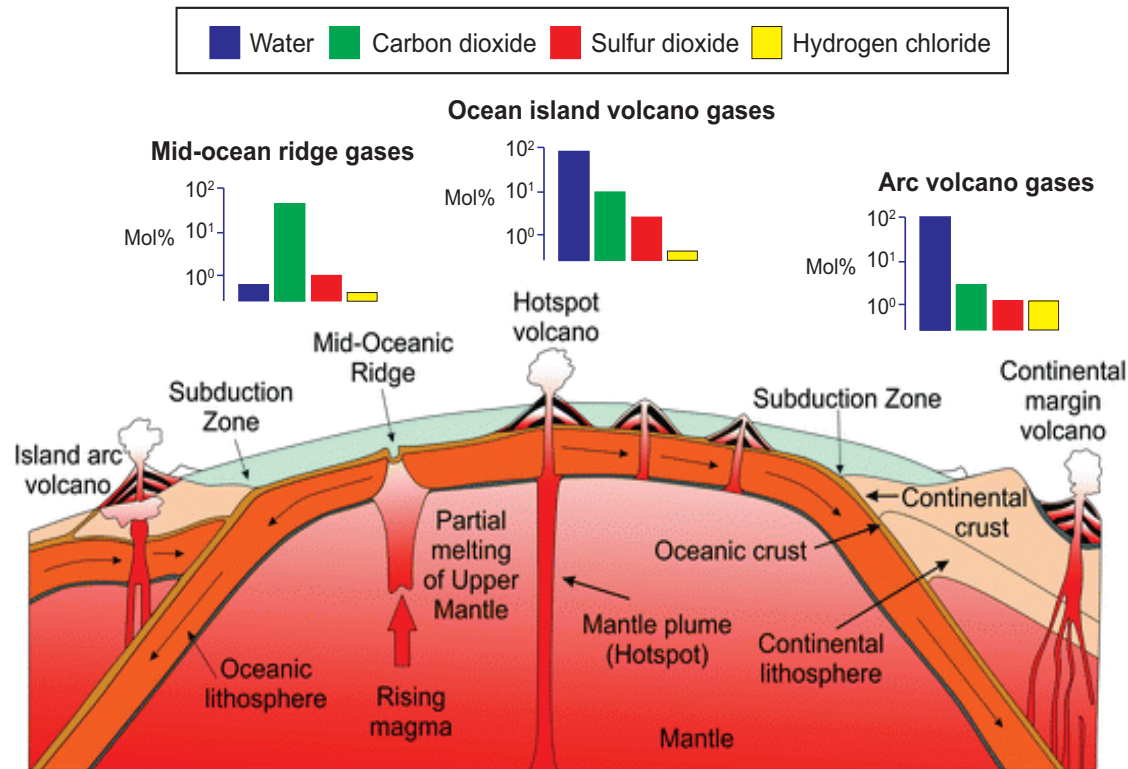


+ Analog experiments and numerical modelling

ISTO, Orléans, CNRS

# VOLCANIC GASES AND GEODYNAMICS

## chemical variations and sources (isotopes)



(e.g. Anderson 1975; Gerlach, 1980, 1981; Allard 1983, 1986; Taylor 1986; Gerlach and Graeber, 1985; Symonds et al., 1994; Giggenbach 1996; Hilton et al 2002; Fisher 2008; Métrich and Wallace, 2008; Edmonds et al 2015; Aiuppa et al., 2016; Shinohara 2018)



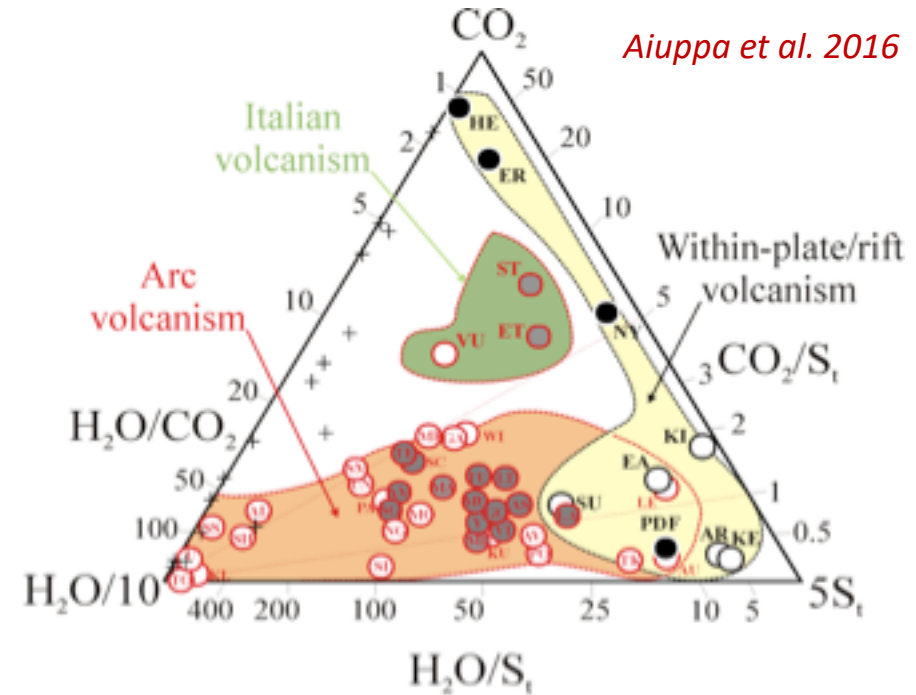
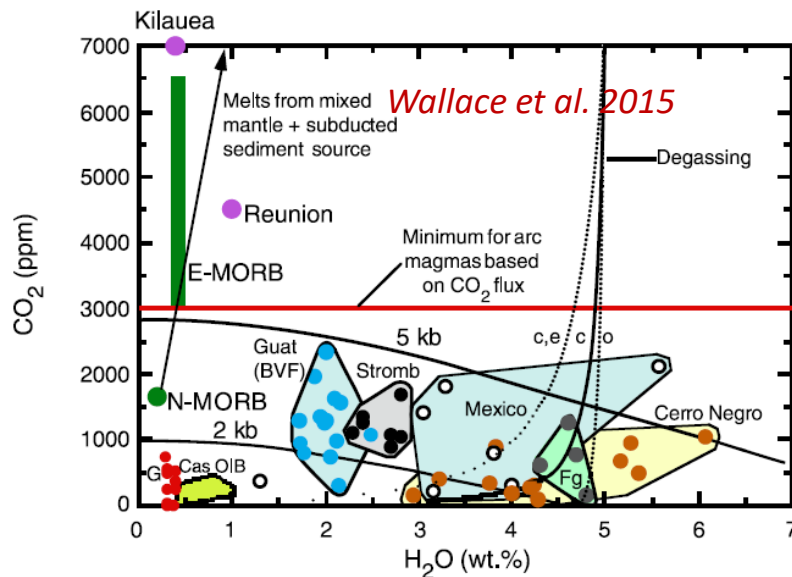
# Volcanic gases in different tectonic settings have different compositions, in good agreement with the dissolved volatile record in magmas

## Arc volcanic gases

- rich in water
- $C/S < 10$  (but mostly between 1 and 5)

## Non-arc

- poorer in water
- wider spread in  $C/S$



Consistent with the volatile record in **melt inclusions** trapped in crystals and, hence, with **volatile abundances in magma sources** !

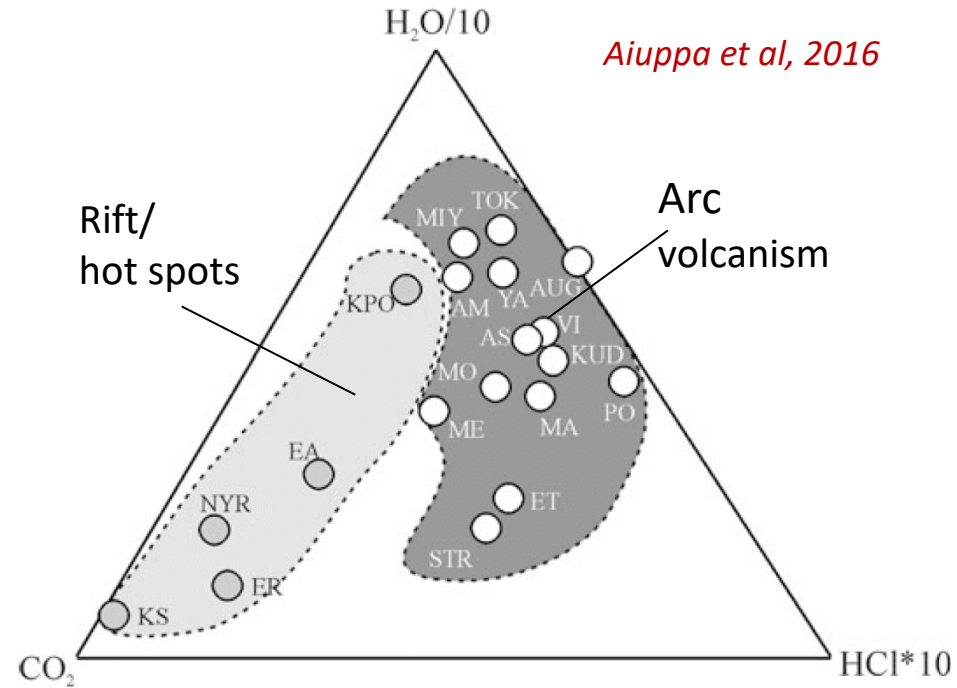
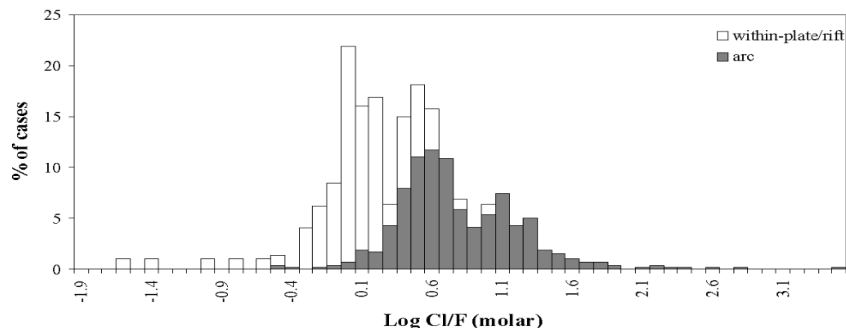
# Volcanic gases in different tectonic settings have different compositions, in good agreement with the dissolved volatile record in magmas

## Arc volcanic gases

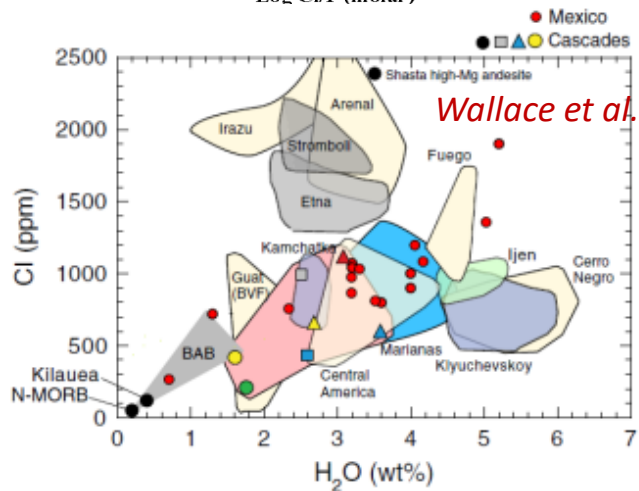
- richer in both water and chlorine
- Cl/F ratio systematically higher

## Non-arc

- richer in both CO<sub>2</sub> and fluorine



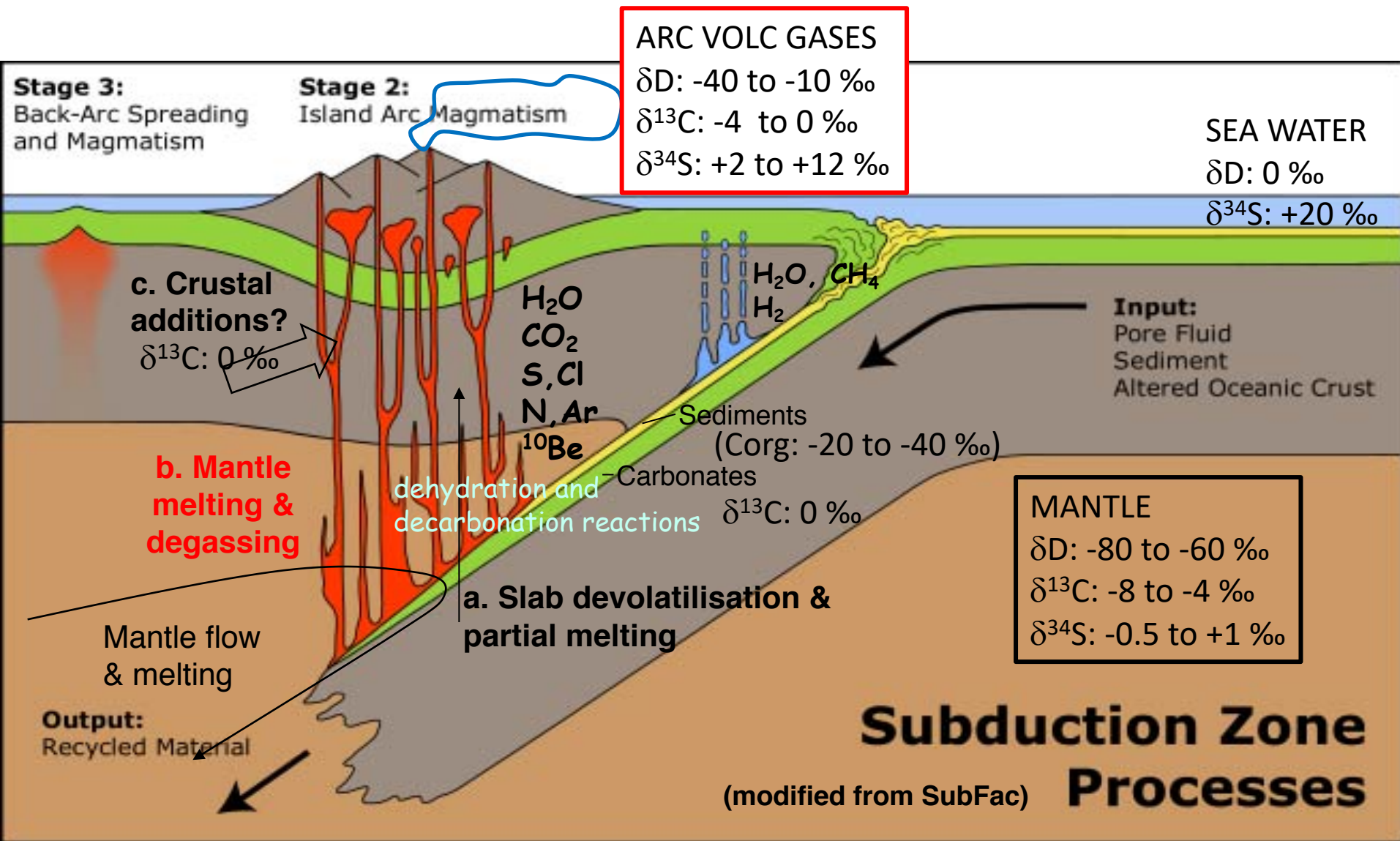
*Aiuppa et al, 2016*



*Wallace et al. 2015*

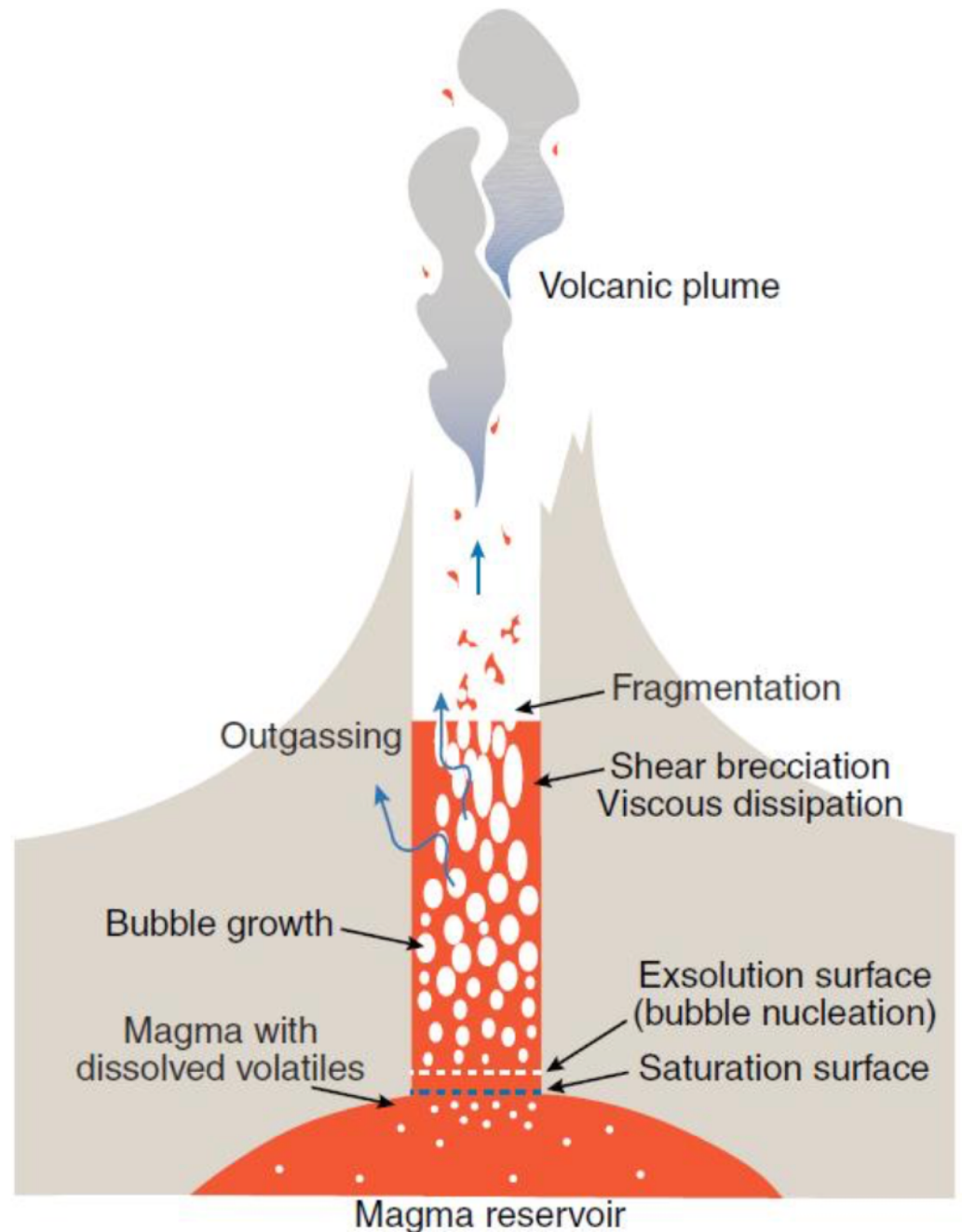
Again, consistent with the volatile record in melt inclusions trapped in crystals and, hence, with volatile abundances in magma source regions !

Isotopic tracers confirm that arc volcanic gases are enriched in volatiles (H<sub>2</sub>O, C, S, N) derived from the subducting plates (± crustal additions)



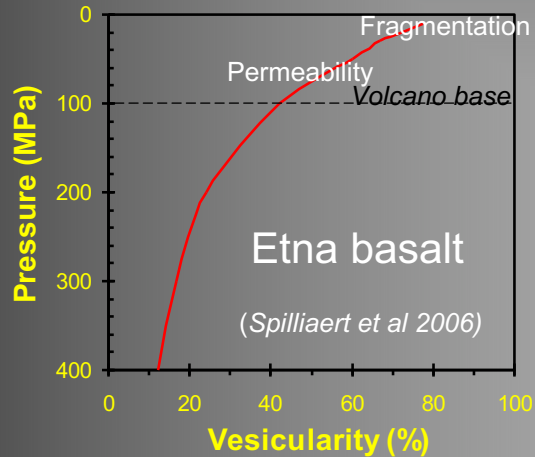
# VOLCANIC GASES AND ERUPTION PROCESSES

## Gas survey and precursors



# Key properties and roles of the magmatic gas phase in magma dynamics

## Magma



Crystallisation paths

Rheological properties  
( $T_L$ ,  $\rho$ ,  $\mu$ ,  $\beta_m$ )

## Gas phase

**The driving force of eruptions!**  
(control of eruption style & intensity)

**Highly mobile (fast vector of info on underground processes)**  
(bubble expansion & coalescence, foams, separate upflow, gas precursors)

**Composition evolves as pressure decreases (diff. volatile solubilities)**  
 $\text{CO}_2 \ll \text{H}_2\text{O} < \text{S} < \text{Cl} < \text{F}$

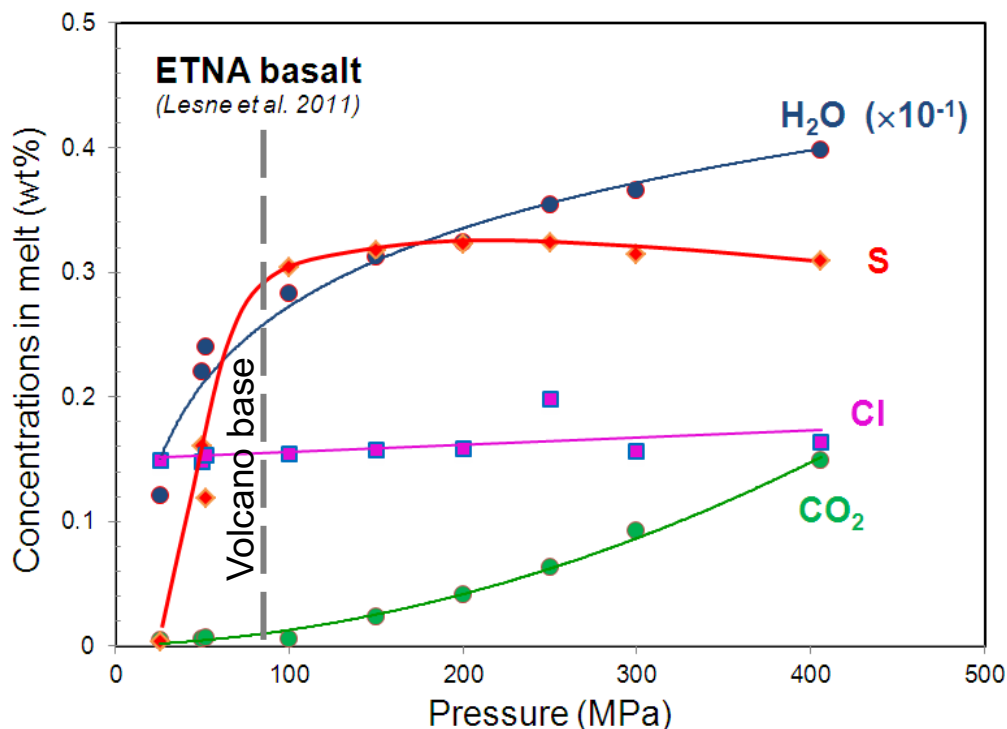
**Bubble nucleation & growth**  
(volatile exsolution, diffusion/expansion)

Decompression of 1 m<sup>3</sup> of melt containing 1 wt% dissolved water at 1 kbar (3 km) generates about 100 m<sup>3</sup> of gas at 1 bar and 1000°C

# Multi-component (H<sub>2</sub>O, CO<sub>2</sub>, S, Cl) solubility evolutions during magma decompression (HT-HP lab experiments and crystal melt inclusions)

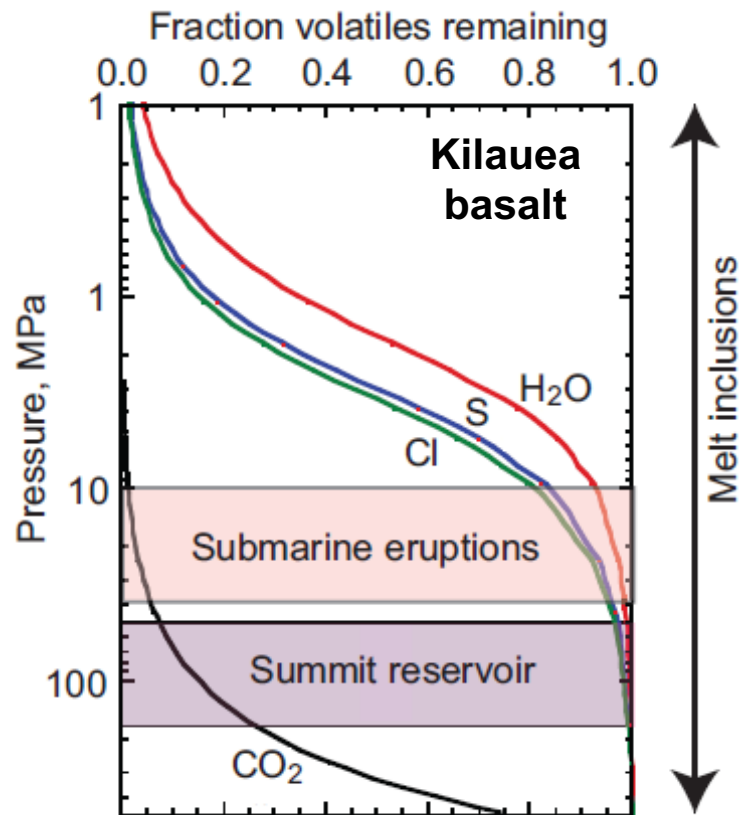
General exsolution order and bulk loss rate: **CO<sub>2</sub> >> H<sub>2</sub>O ≥ S > Cl > F**

## Experimental measurements



At volatile saturation

$$P_{\text{CO}_2} + P_{\text{H}_2\text{O}} \cong P_{\text{tot. fluid}} \cong P_{\text{lithostatic}}$$



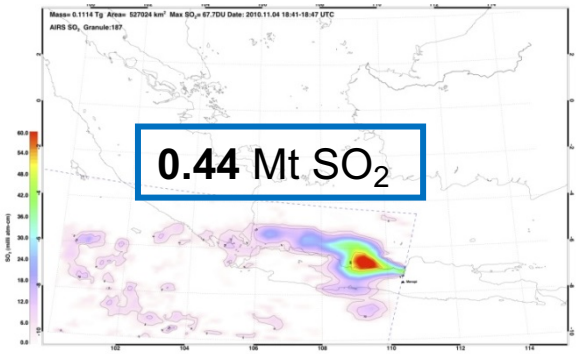
Edmonds et al. 2014, and ref. therein

**CO<sub>2</sub>: the first volatile species to form gas bubbles in melts at depth and, therefore, a key tracer of deep gas supply**

# MERAPI: CO<sub>2</sub> precursor of the November 2010 sub-Plinian eruption

(~ 1.5 million evacuees, ~400 casualties)

Intermittent gas sampling

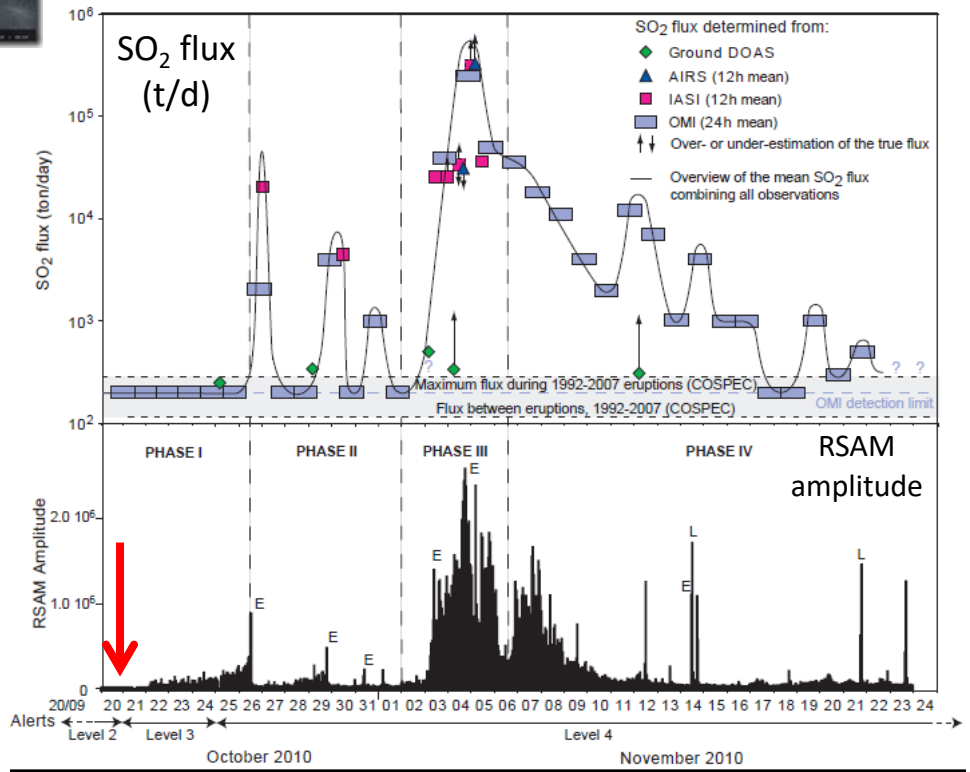


Surono et al. JVGR 2012

2010 Fumarole gas analyses (mol%)

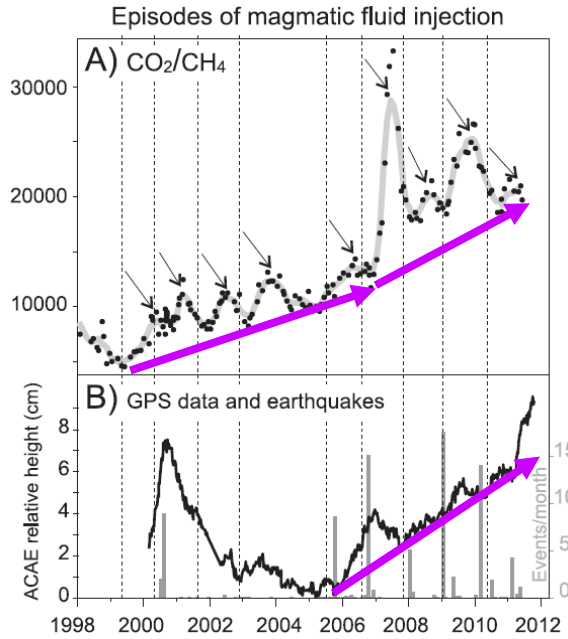
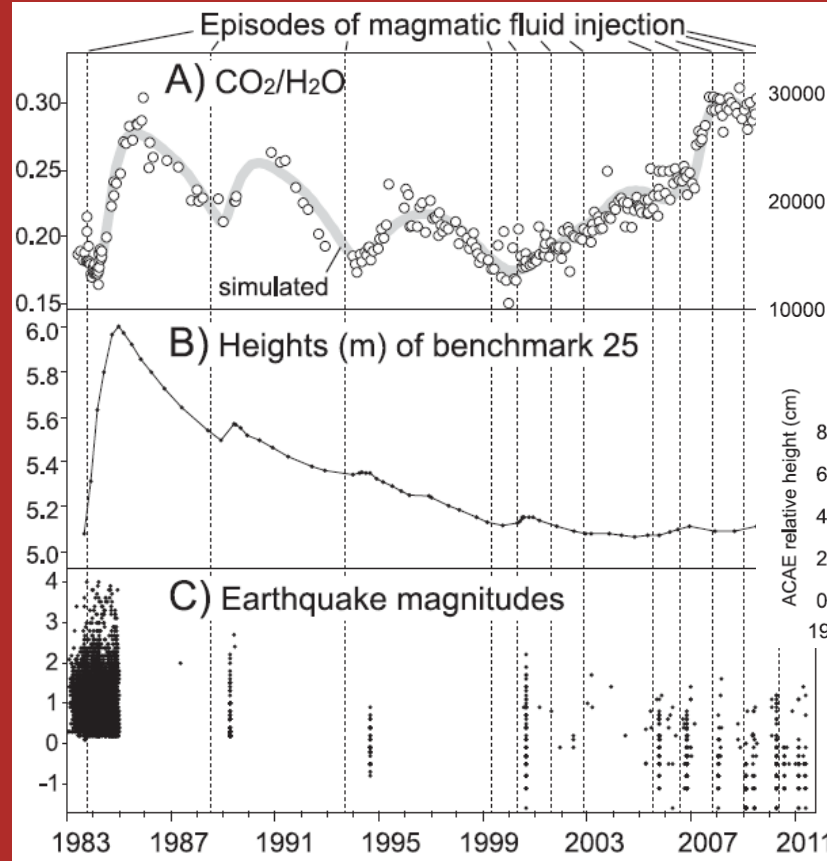
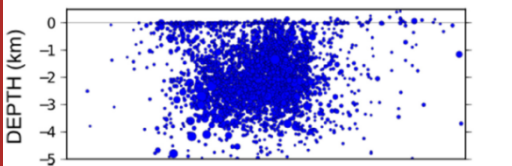
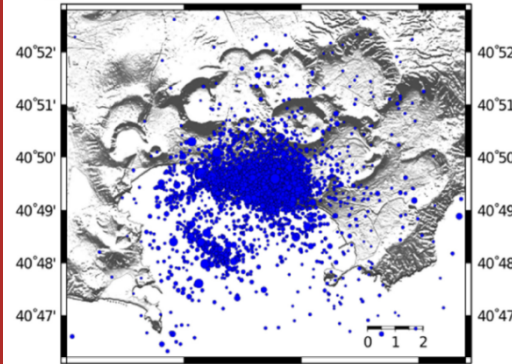
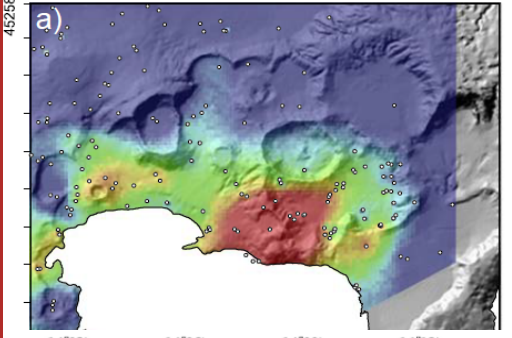
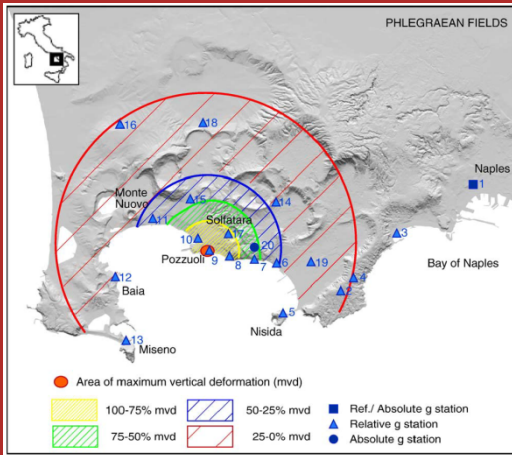
	26 May	Sept. avg	20-Oct.
T (°C)	460	575	575
H <sub>2</sub> +O <sub>2</sub>	0.07	0.0013 <sup>c</sup>	0.02 <sup>c</sup>
N <sub>2</sub>	1.1	0.1	0.02
CH <sub>4</sub>	0.01	n.d.	0.01
CO	n.d.	0.01	0.03
CO <sub>2</sub>	5.6	10	34.6
SO <sub>2</sub>	0.8	1.0	0.3
H <sub>2</sub> S	0.2	0.45	2.5
HCl	0.2	0.36	0.6
HF	n.d.	n.d.	n.d.
NH <sub>3</sub>	0.01	0.5	2.8
H <sub>2</sub> O	92	87	58.8

CO <sub>2</sub> /S	5.6	6.3	12
CO <sub>2</sub> /HCl	28	28	58
CO <sub>2</sub> /H <sub>2</sub> O	0.06	0.1	0.6



# Fumarolic gas survey - Long-term signals

## Ex.: stepwise inputs of magma-derived CO<sub>2</sub> and caldera unrest at Campi Flegrei, Italy



*Chiodini et al. 2012*



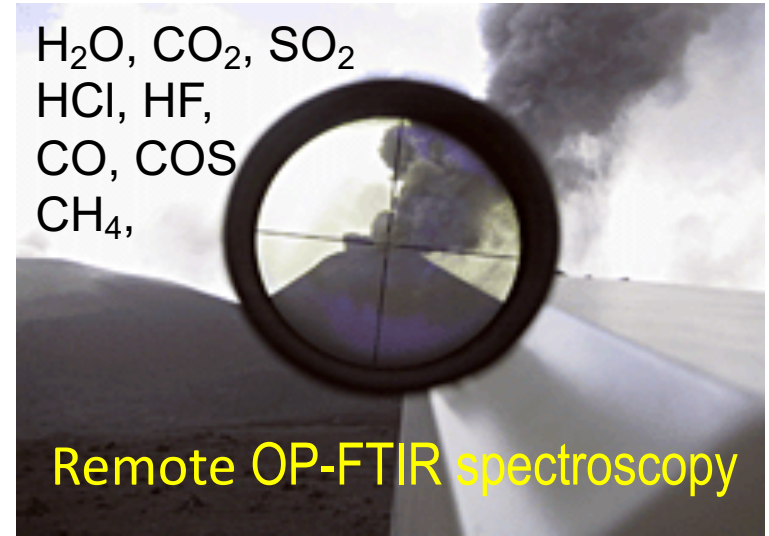
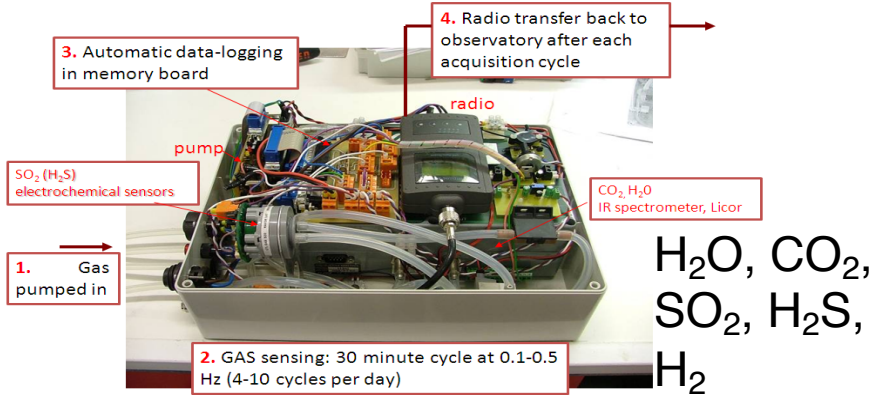
*Routine gas sampling*



# A revolution in last two decades: new tools allowing high-frequency ( $\geq 1$ Hz) in situ or remote sensing of volcanic gas composition, even during eruptions!

## In situ MultiGAS plume analysis

In situ method: gas is pumped in and measured at sensors in 4 fully automated STEPS



Etna - Italy



Miyakejima (Japan)



Shinohara et al. 2005; Aiuppa et al. 2005, 2011, 2016; De Moor et al. 2017, Moussallam et al. 2018...



Mori et al. 1993, 1995; Francis et al. 1998; Burton et al. 2000, 2007; Oppenheimer et al 2008; Sawyer et al., 2008a-b; Edmonds & Gerlach 2009; Allard et al. 2002, 2005, 2012, 2016

Bulk gas composition scaled to  $\text{SO}_2$  flux  $\Rightarrow$  total gas flux

# Coupled with UV remote quantification of volcanic SO<sub>2</sub> fluxes

(absorption of scattered sunlight by SO<sub>2</sub> in the 300-325 nm UV band)

## COSPEC (Correlation Spectrometry)

Stoiber & Malone 1973  
Stoiber et al. 1983

UV

(used in mobile or fixed scanning modes)

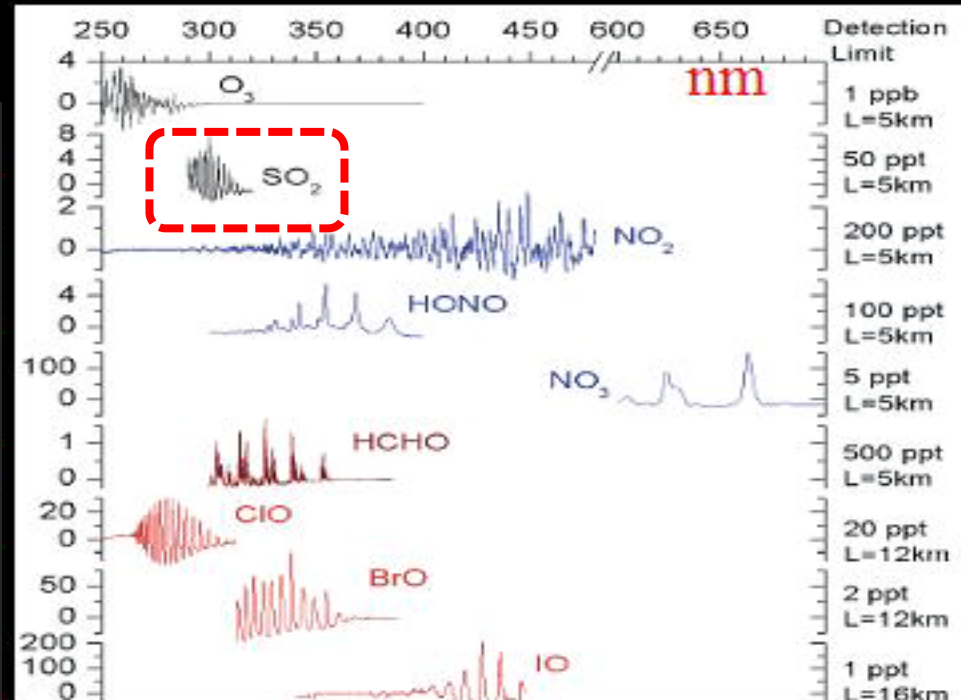
## Mini-DOAS (Differential Optical Absorption Spectroscopy)

Galle et al. 2000  
Arellano et al. 2012

## Dual UV-cameras

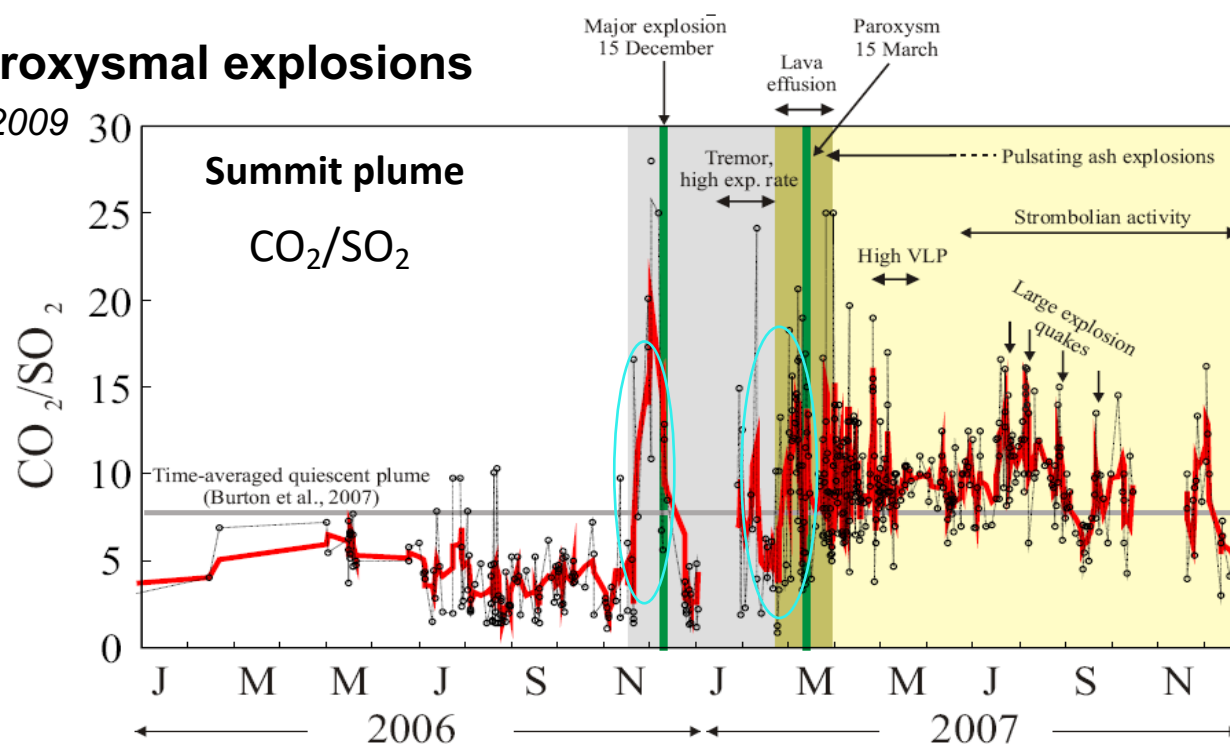
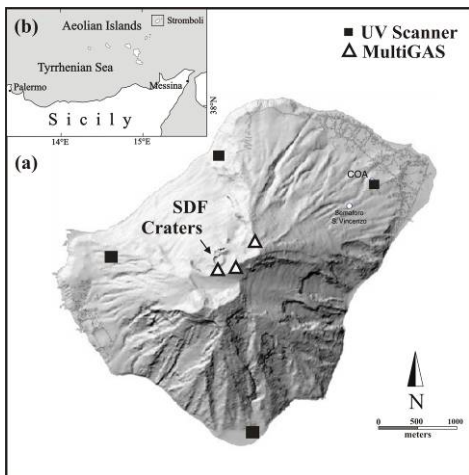
(high-resolution SO<sub>2</sub> imaging, up to 25 Hz)

Mori & Burton 2005  
Kern et al. 2008  
Tamburello et al. 2013



# CO<sub>2</sub> precursors of paroxysmal explosions

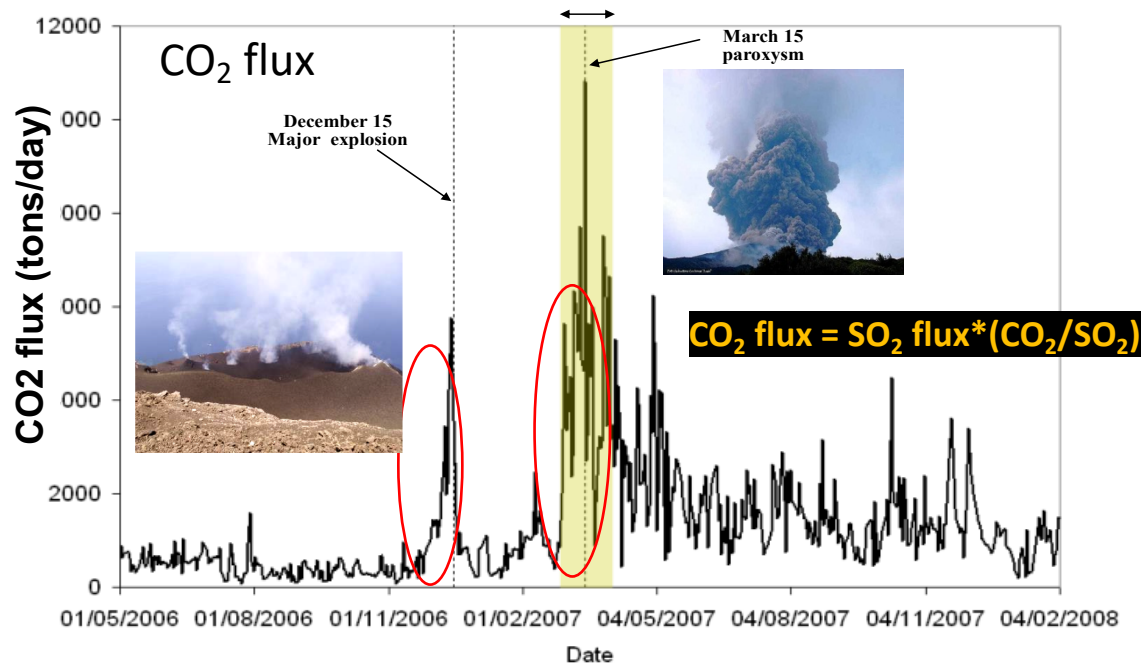
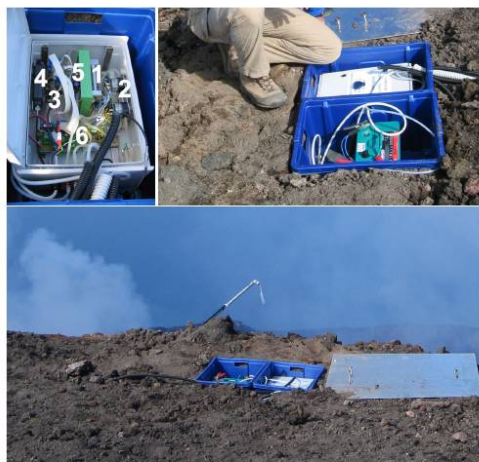
**Stromboli: Aiuppa et al. 2009**



## Multi-GAS

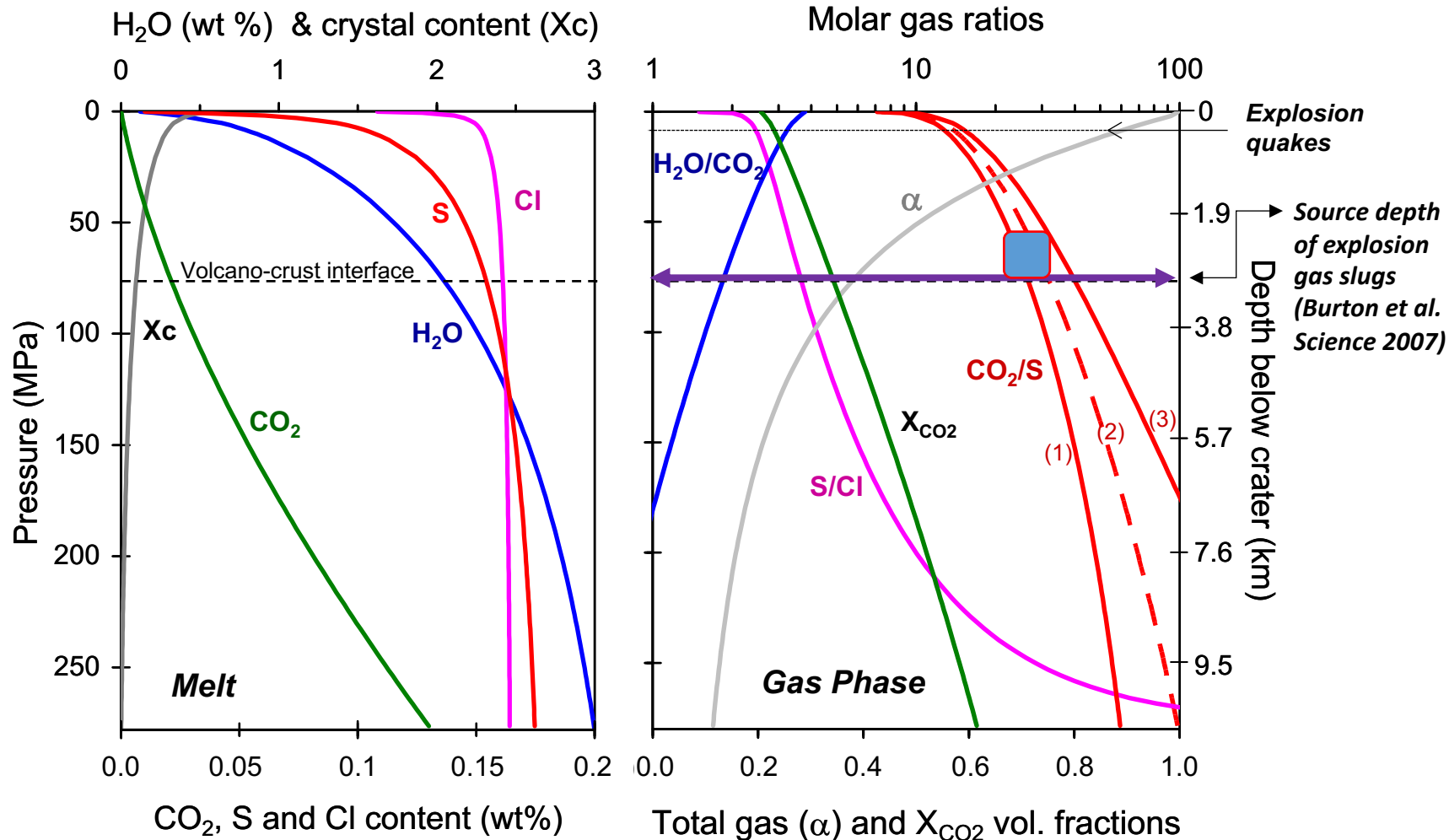
CO<sub>2</sub>, H<sub>2</sub>O: IR spectrometry

SO<sub>2</sub>, H<sub>2</sub>S: electrochemical sensors



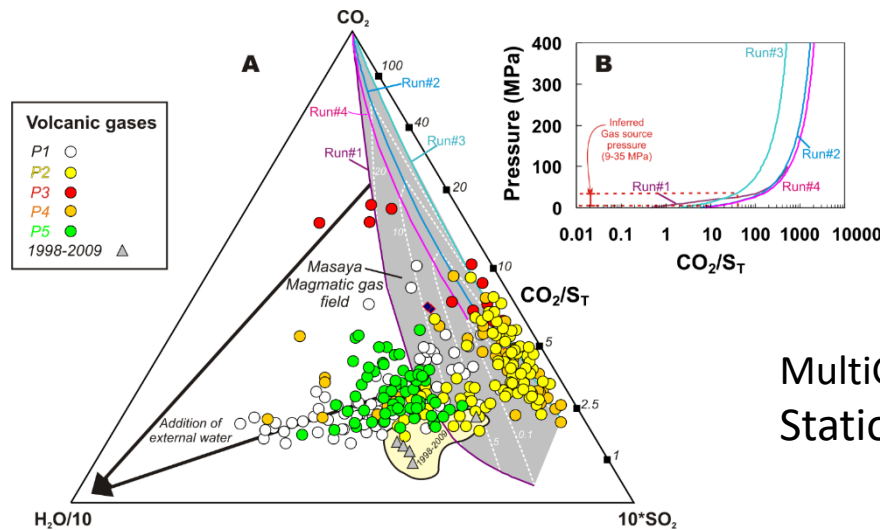
# Volcanic gas ratios as 'geobarometers'

Ex.: *P*-related evolution of molar ratios in the magmatic gas phase during CSD decompression of Stromboli basalt from 280 MPa (~10.5 km depth) to the surface, computed from the measured amounts of dissolved volatiles in the melt (*Mis*) (Allard, 2010)



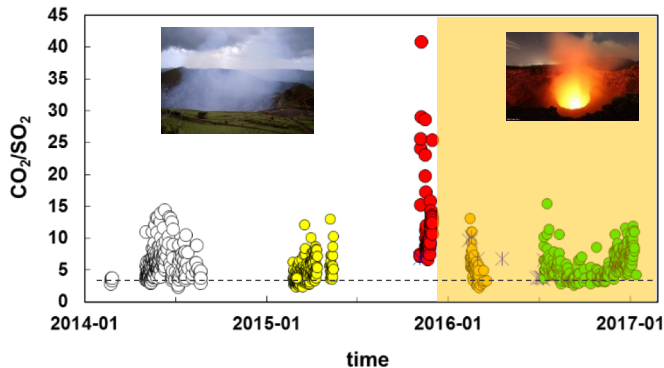
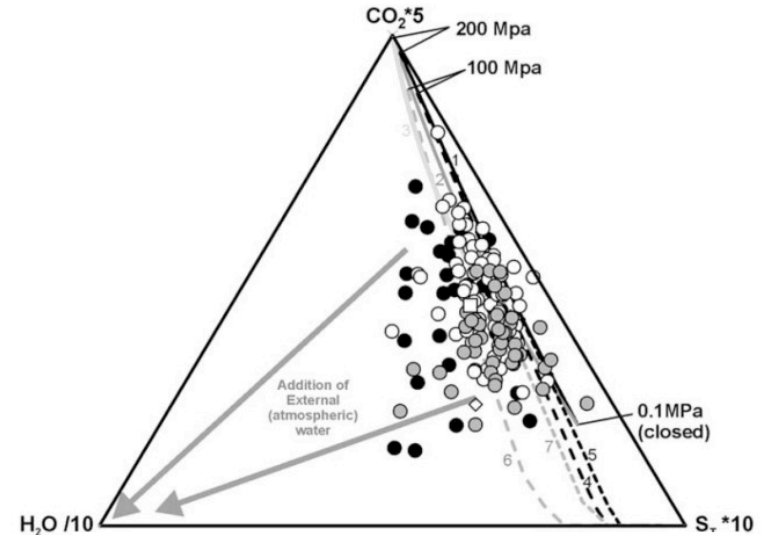
Also, solubility and thermodynamic models (Newman & Lowenstern 2002; Papale et al., 2006; Burgisser and Scaillet 2008; Moretti et al., 2013....)

# Masaya, Nicaragua (2014-2017)

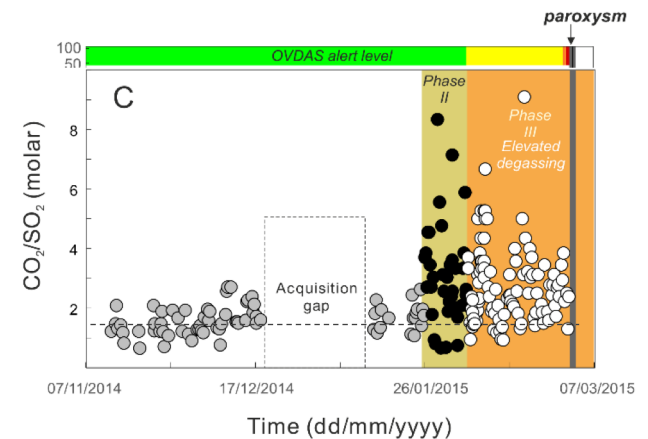


MultiGAS Stations

# Villarrica, Chile (2014-2015)



Aiuppa et al. 2016  
Aiuppa et al. 2018

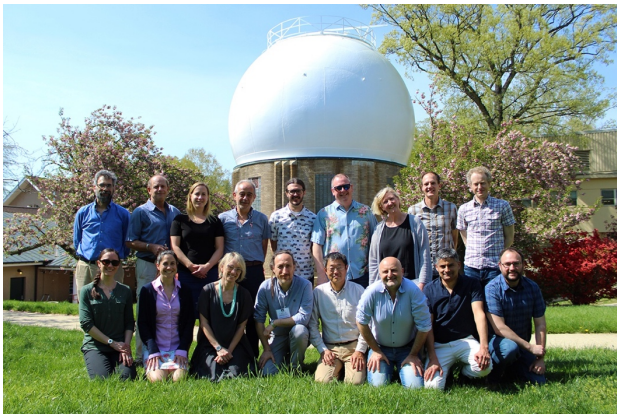


**$CO_2/SO_2$  ratio variations weeks/months prior to unrest or eruption due to supply of  $CO_2$ -rich, deeply sourced (10-40 MPa) bubbles during magma decompression**

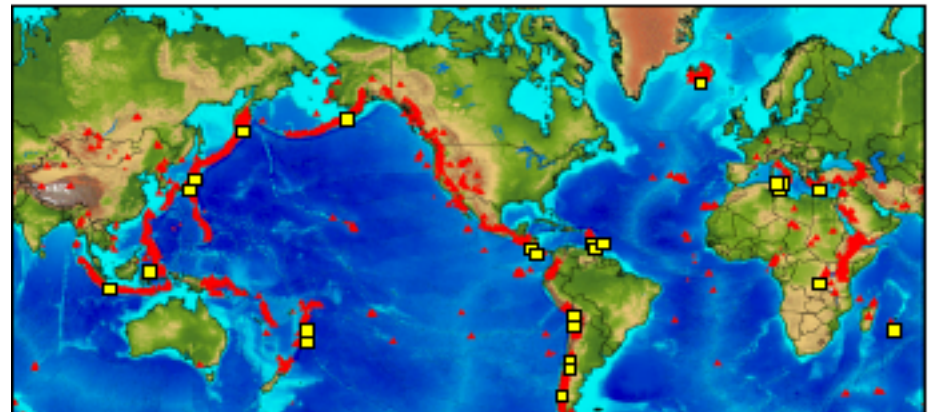
## The DCO-DECADE (DEep CARbon DEgassing) Initiative, 2012-2019

### 3 main goals

- A. Improve **current estimates of deep carbon emission budget from global subaerial volcanism and active lithospheric regions**, in particular from subduction zones.
- B. Develop a **network for continuous CO<sub>2</sub> survey on about 25-30 of the most actively degassing volcanoes on Earth**, in connection with volcano Observatories & Agencies.
- C. Build up a **database** for global deep carbon emissions from volcanic and lithospheric regions (plumes, hydrothermal fluids, soil emanations, groundwater flows, etc.)



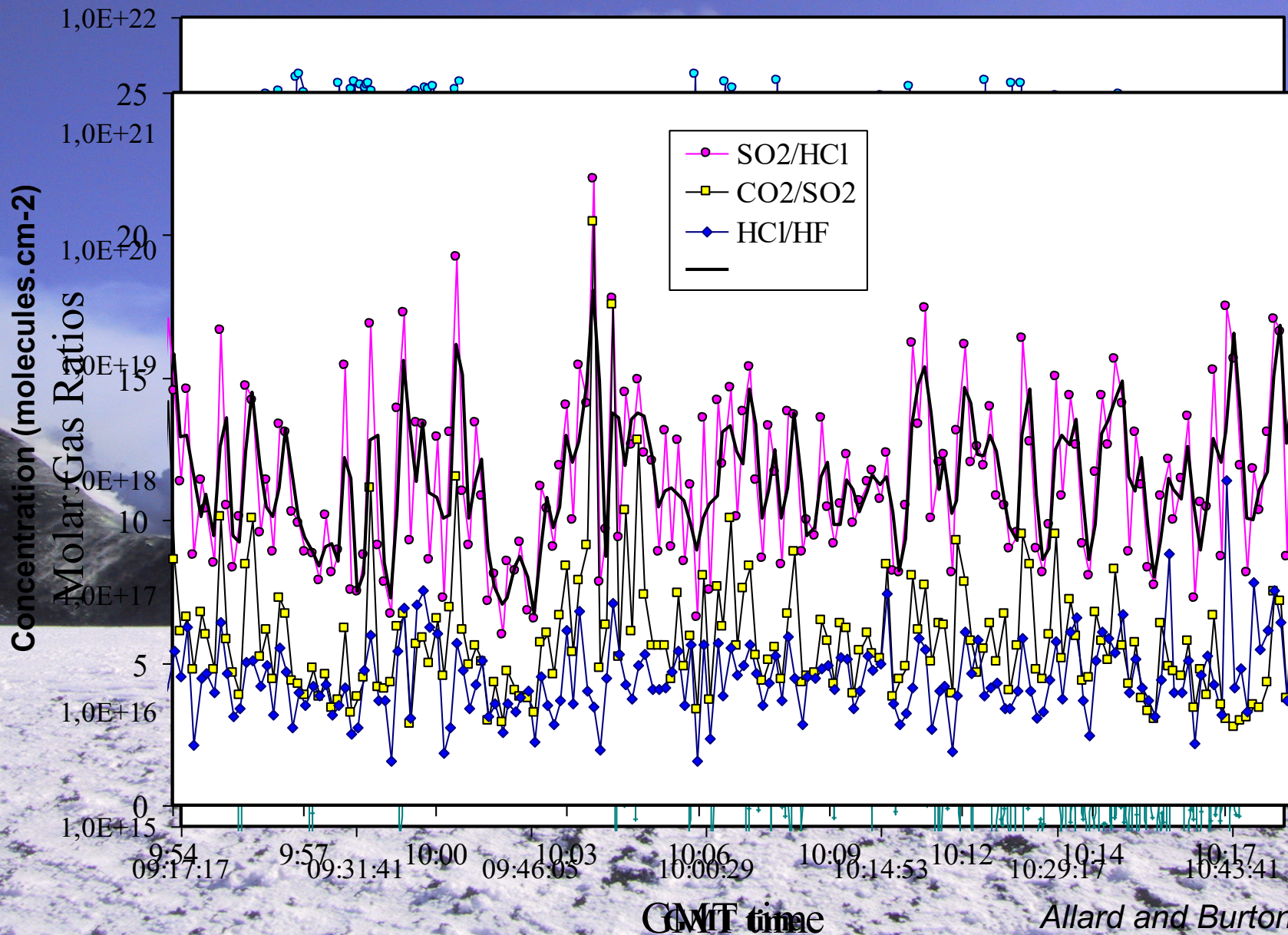
Board Synthesis Meeting, Carnegie Washington, April 2018



28 main volcanoes worldwide now permanently monitored with **MultiGAS stations**

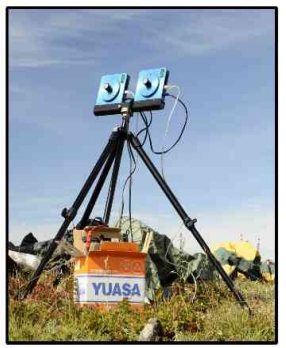
# HF OP-FTIR sensing: example of explosive degassing at Etna 14/12/2002

d = 400 m, 1 FTIR spectrum every 4 sec. = 900 gas samples in one hour!



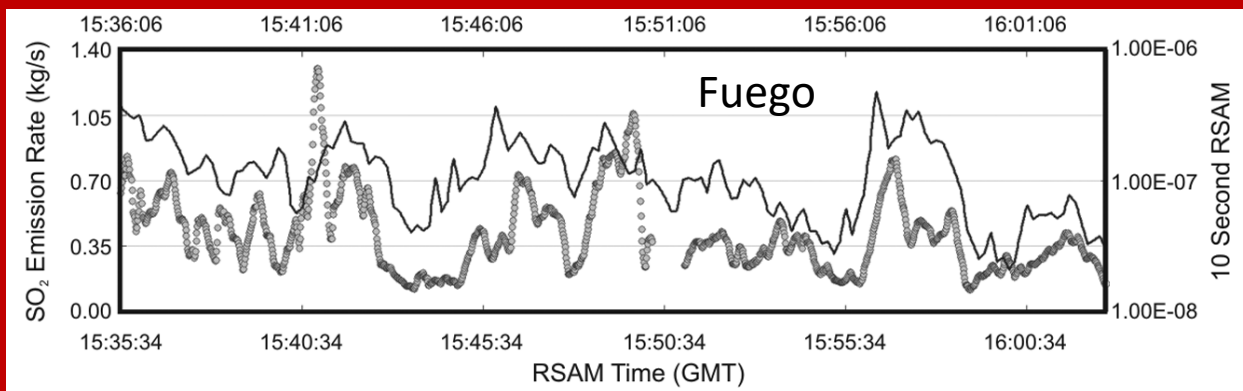
# High frequency $SO_2$ flux measurements

Correlate well with shallow-sourced geophysical signals  
(tremor, LP-VLP)



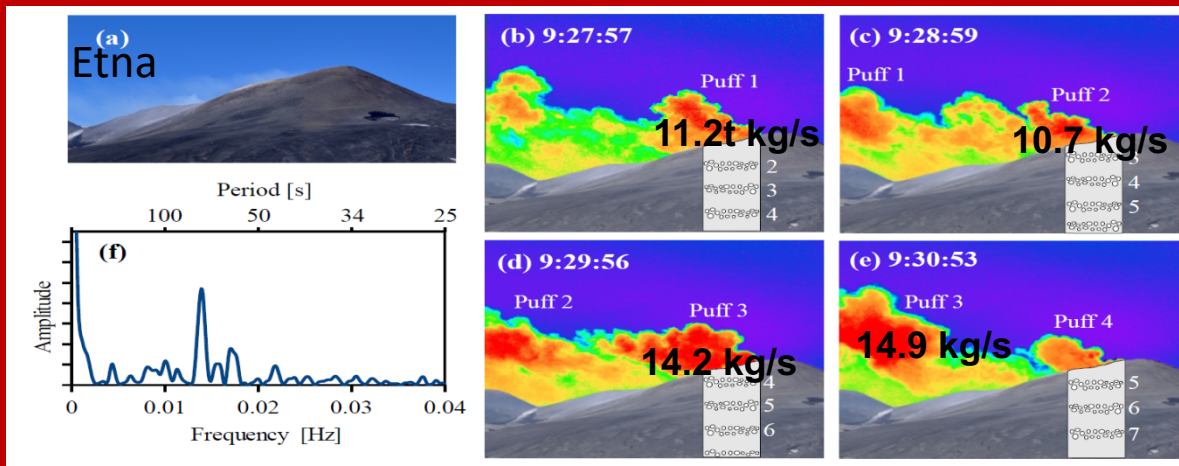
## Dual UV-Camera imaging: 1-5 Hz $SO_2$ flux time-series

- *FUEGO (Guatemala): co-variations of  $SO_2$  flux and RSAM (Nadeau et al. 2011)*



- *ETNA: Distinct periodic structure in conduit bubble layering (Tamburello et al. GRL 2013)*

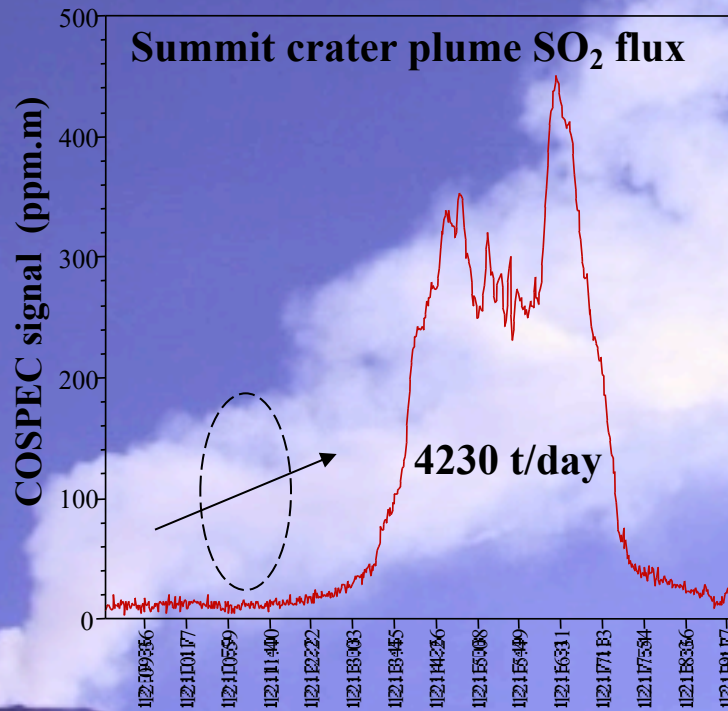
2 band-pass optical filters:  
- 310 nm ( $SO_2$ )  
- 330 nm (aerosols & particles)





# Volcanic SO<sub>2</sub> fluxes and magma degassing rates (by scaling to the magma sulfur content)

$$V_d = \frac{1}{2} F_{\text{SO}_2} / [\Delta S * \rho_m * (1 - X_c)]$$



$\Delta S \cong$  initial magma S content

$\rho_m$  = magma density

$X_c$  = crystal vol. fraction

$(1 - X_c)$  = melt vol. fraction

$V_d/V_e$  ratio:

total degassed vs erupted magma

Mount Etna

# EXCESS DEGASSING: most volcanoes actually emit more to much more gas than allowed by co-erupted magma volumes !

## ❖ Persistently degassing open-conduit volcanoes



## Examples (time-averaged $V_d/V_e$ )

- Etna: 4 (Allard 1997; Allard et al. 2006)
- Merapi: 7 (Allard et al. 2011)
- Villarica: 9 (Witter et al. 2005)
- Stromboli: 10 (Allard et al. 2008)
- Ambrym: 16 (Allard et al. 2016)
- Yasur: 18 (Métrich et al. 2011)
- Lava lakes:  $>10^2$ - $10^4$

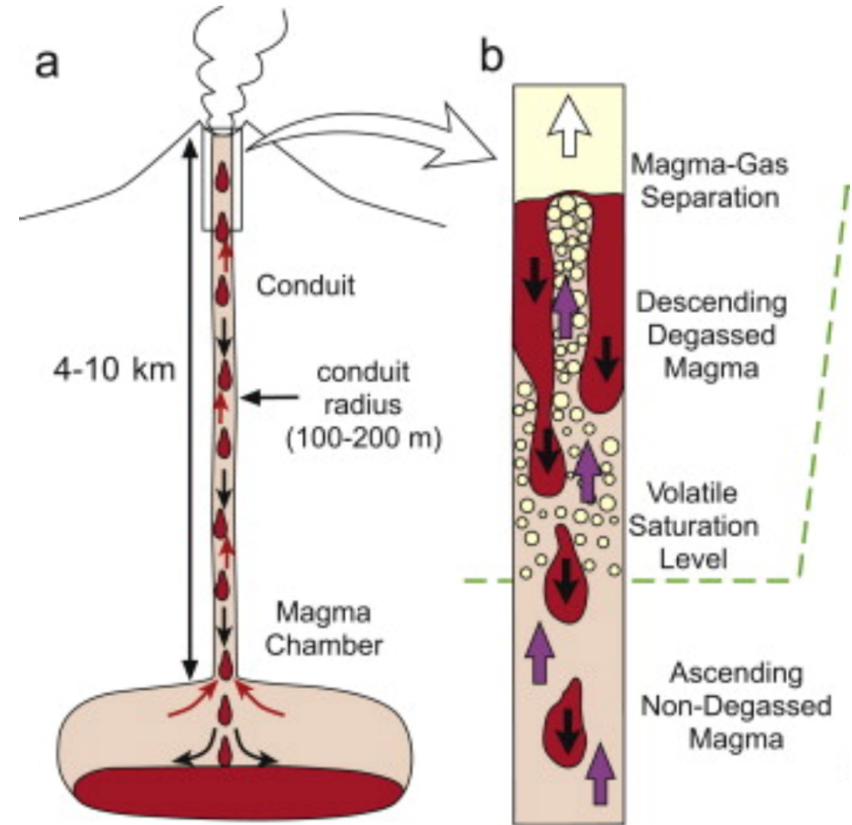
## **Probable mechanisms (single or combined):**

- *Differential bubble transfer across volcanic systems and conduits*
- *Gas percolation through permeable magma or/and sheared stress conduit walls*
- *Convective magma overturn in conduit-reservoir systems*
- **Major IMPLICATION: growth of large bodies of degassed (solidified) magma beneath many volcanoes (e.g. at Etna, Asama, etc....)**

**Nyiragongo, Dem. Rep. Congo**



**Top ten volcanic CO<sub>2</sub> emitters**

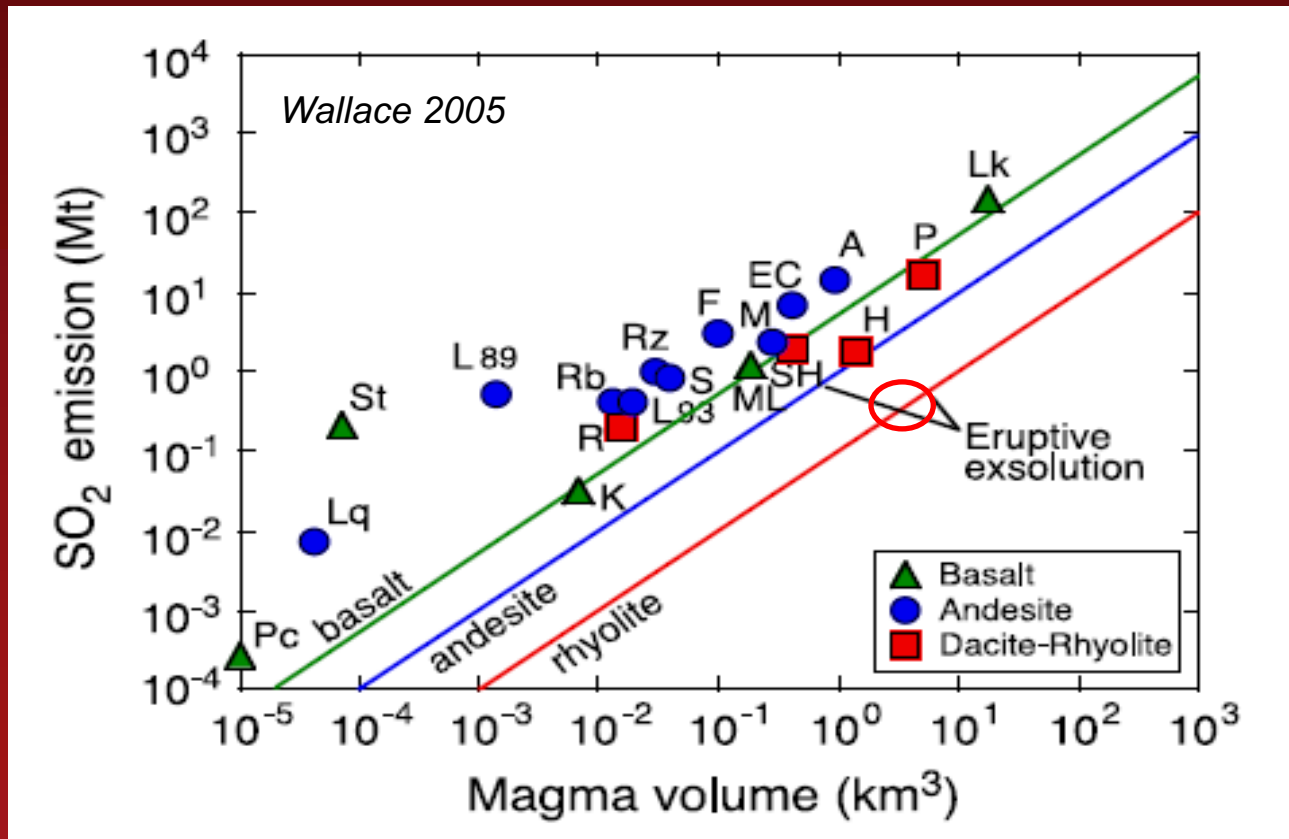


Shinohara, 2008  
Burton et al., 2013

Volcano	Country	CO <sub>2</sub> Flux (t/d)	CO <sub>2</sub> Flux (Mt/yr)
Nyiragongo	DR Congo	52,410	19.13
Popocatepetl	Mexico	29,000	10.59
Ambrym	Vanuatu	20,000	7.30
Etna	Italy	16,363	5.97
Miyakejima	Japan	14,500	5.29
Oldoinyo Lengai	Tanzania	6,630	2.42
Kīlauea	USA	6,549	2.39
Stromboli	Italy	1,991	0.73
Masaya	Nicaragua	1,935	0.71
White Island	New Zealand	1,780	0.65
Augustine	USA	1,760	0.64
Erebus	Antarctica	1,630	0.59
Soufrière Hills	Montserrat	1,468	0.54
Galeras	Colombia	1,020	0.37

Convective magma overturn likely a key process at top ten degassing volcanoes !

❖ Excess degassing also verified during discrete eruptions of various magma types !

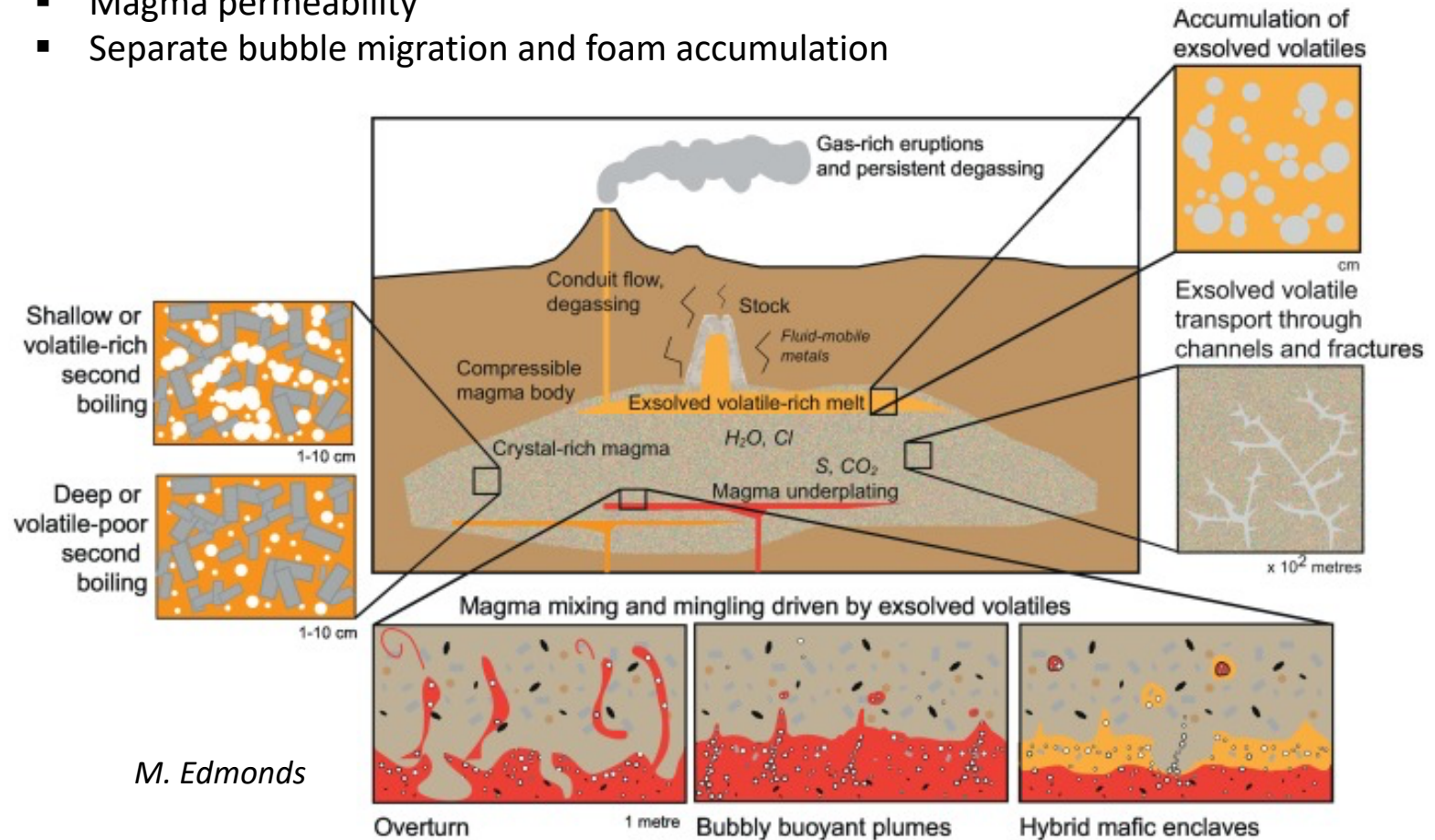


➤ Likely: pre-eruptive accumulation of an exsolved gas phase

*e.g. Pinatubo (Westrich and Gerlach 1992)*

# Hence, multiple complex degassing processes affect volcano feeding systems, deserving further investigations

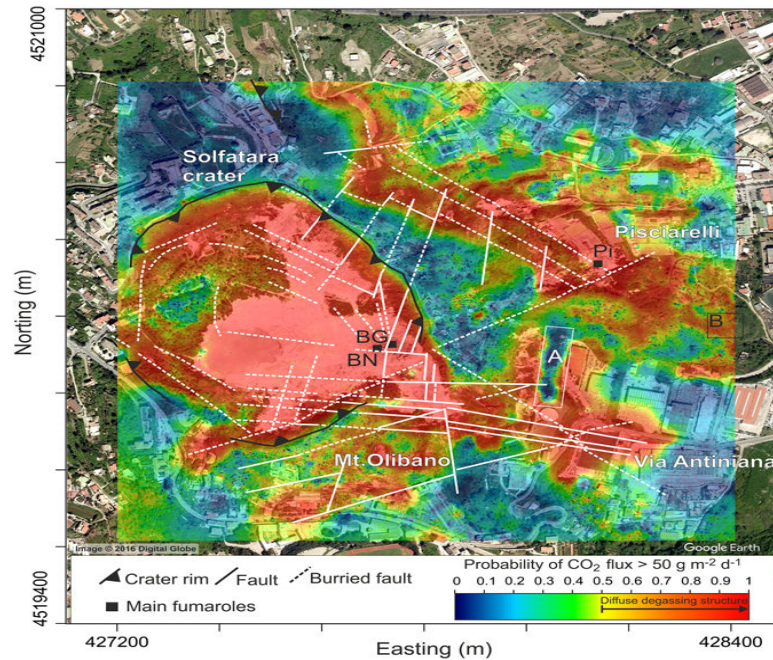
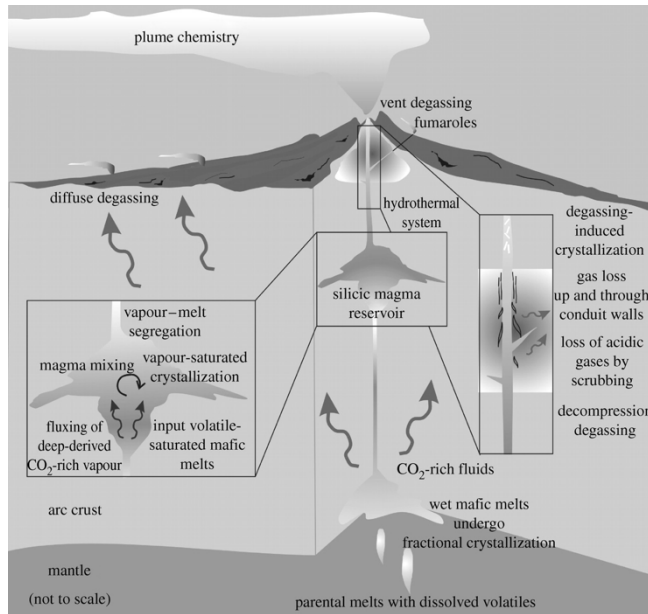
- Gas percolation in crustal magma reservoir (*Bachmann-Bergantz 2006*)
- Magma permeability
- Separate bubble migration and foam accumulation



M. Edmonds

# Another main discovery: diffuse soil degassing through volcanic systems

non-thermal (invisible) emanations of carbon dioxide ( $\pm$  minor  $H_2S$ ,  $H_2$ , He,  $^{222}Rn$ )



Campi Flegre caldera  
*Cardellini et al., 2017*

East African Rift  
*Hunt et al., 2017*

Mammoth Mountains  
*Hill et al., 1998*

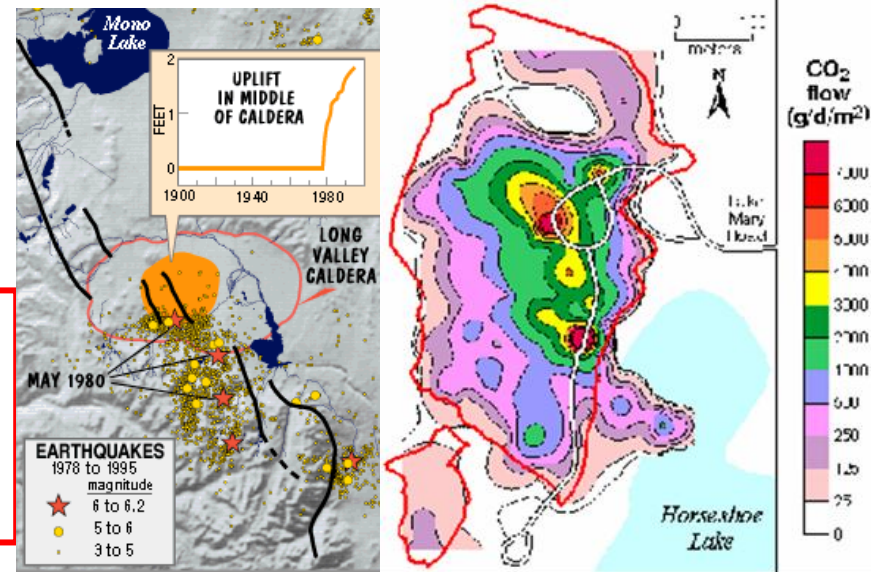


NATURE · VOL 351 · 30 MAY 1991

## Eruptive and diffuse emissions of $CO_2$ from Mount Etna

P. Allard\*, J. Carbonnelle†, D. Dajčević†, J. Le Bronec†,  
P. Morel†, M. C. Robe†, J. M. Maurenas‡,  
R. Faivre-Pierret‡, D. Martin§, J. C. Sabroux\*  
& P. Zettwoog†

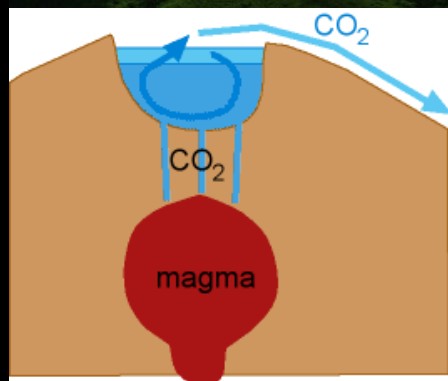
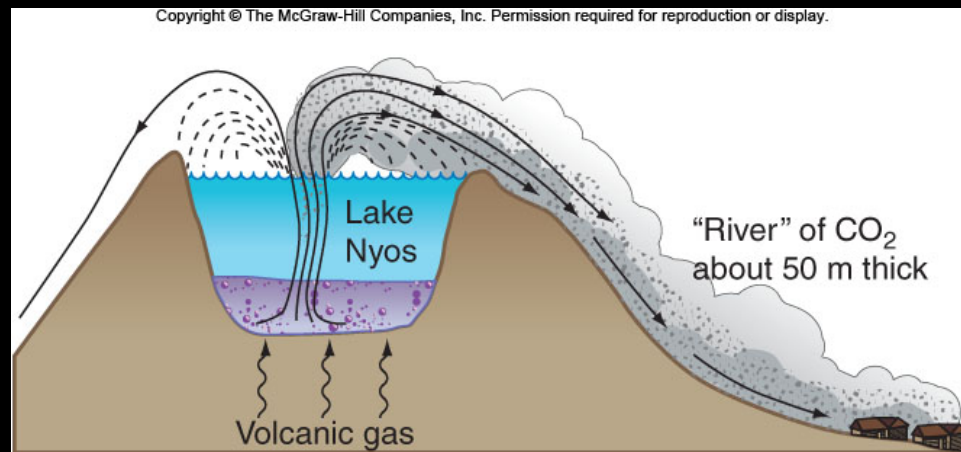
Studying/surveying diffuse volcanic degassing has become a new research field in volcanology over the past 25 years !



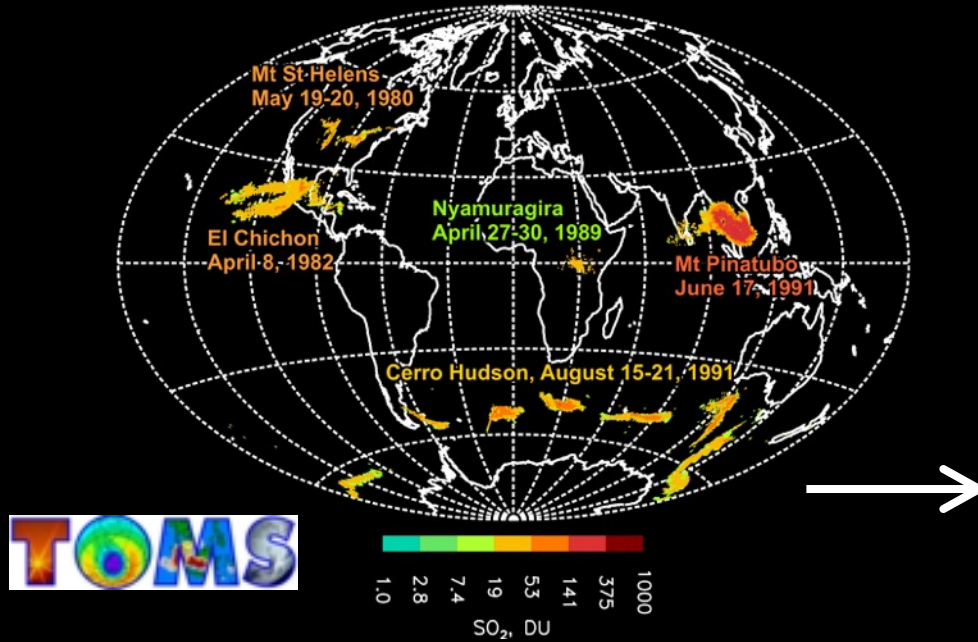
- **Vesuvius 1944:** 8 deaths at the base six months before the eruption
- **Dieng 1979**, Java: 149 persons died while crossing a thick CO<sub>2</sub> stream released during phreatic eruptions (*Le Guern et al 1981; Allard et al 1989*)
- **Lake Nyos 1986**, Cameroon: dense CO<sub>2</sub> flows kills 1700 people up to 16 km distance (*Barberi et al 1987*)

## ❖ THE HAZARD FROM COLD EMISSIONS OF NOXIOUS VOLCANIC CO<sub>2</sub>

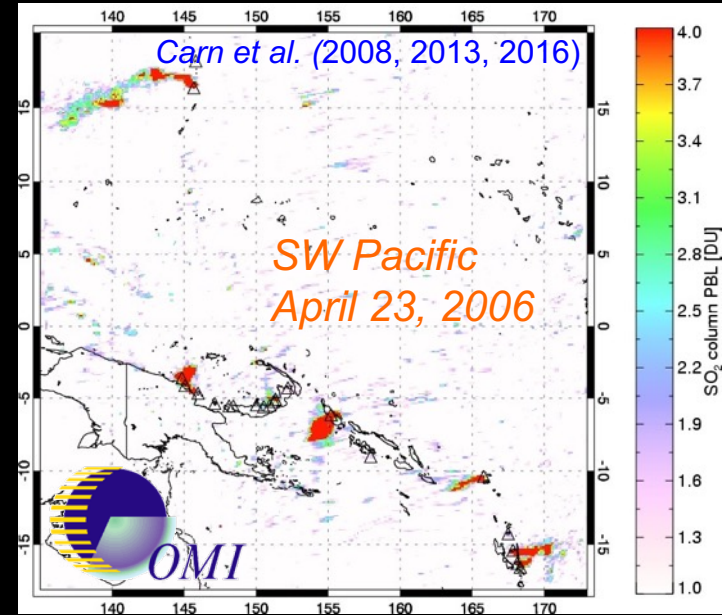
*Gaseous carbon dioxide is 1.6 times denser than air at ambient temperature!*



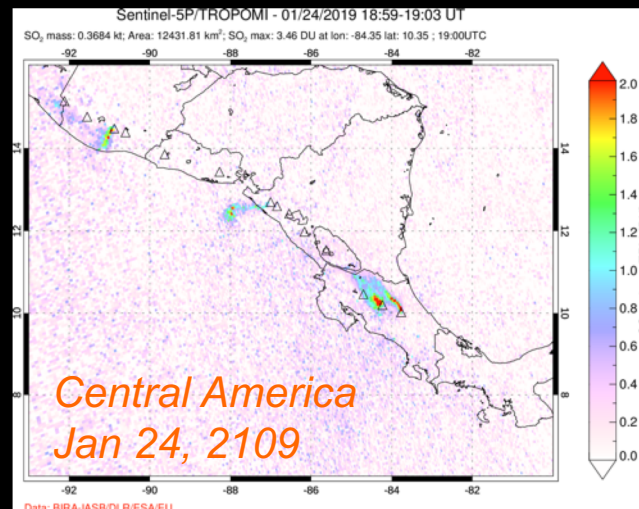
# UV satellite remote sensing of volcanic SO<sub>2</sub>



1978-2005: Total Ozone Mapping Spectrometer (TOMS) -> Eruptive degassing



2004- : Ozone Monitoring Instrument (OMI) -> Eruptive and passive degassing

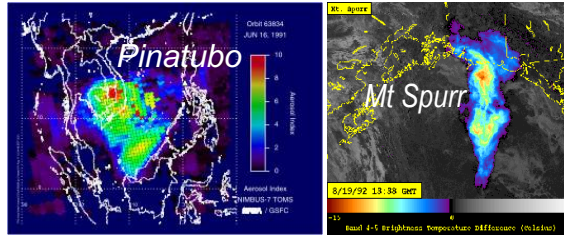
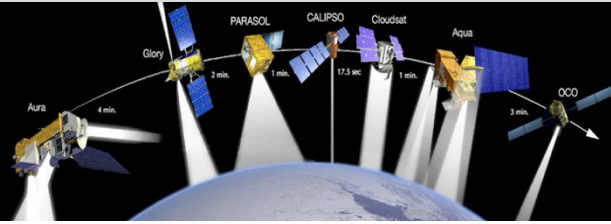


2017- : Sentinel 5P TROPOMI -> Higher spatial resolution of volcanic SO<sub>2</sub> plumes

*Theys et al. (2019)*

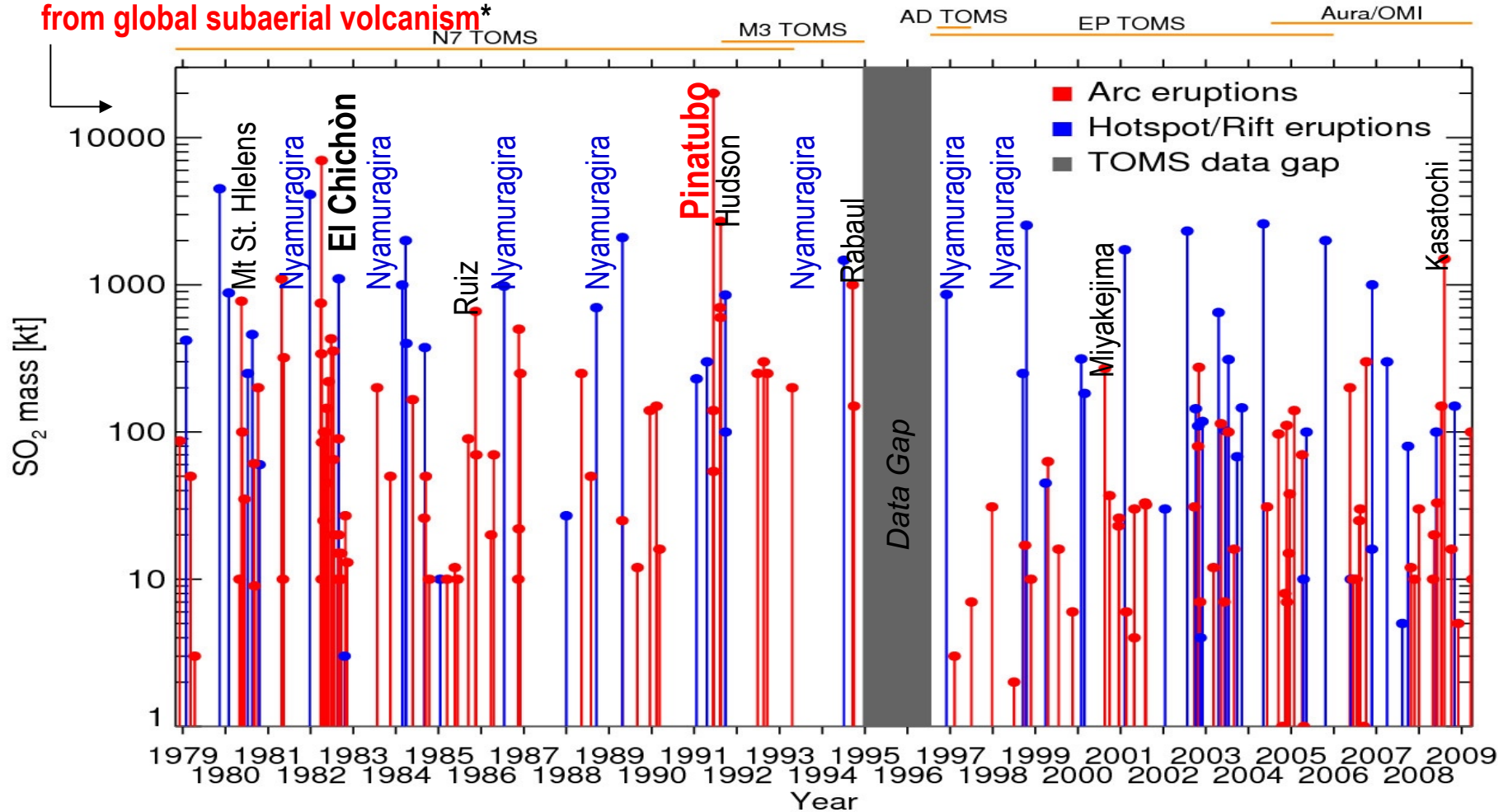


# Space-borne quantification of eruptions' SO<sub>2</sub> mass output



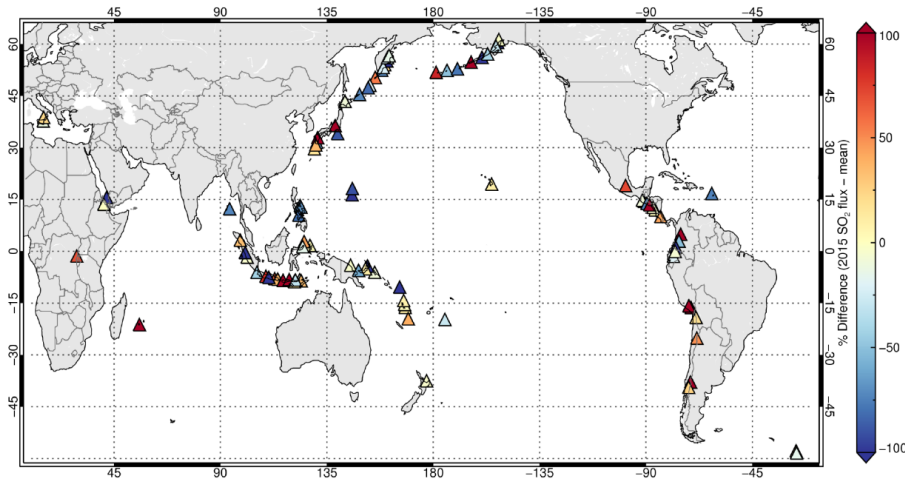
(1979 – 2009, TOMS then OMI)<sup>£</sup>

**Average SO<sub>2</sub> discharge of ~20 Mt/yr from global subaerial volcanism\***

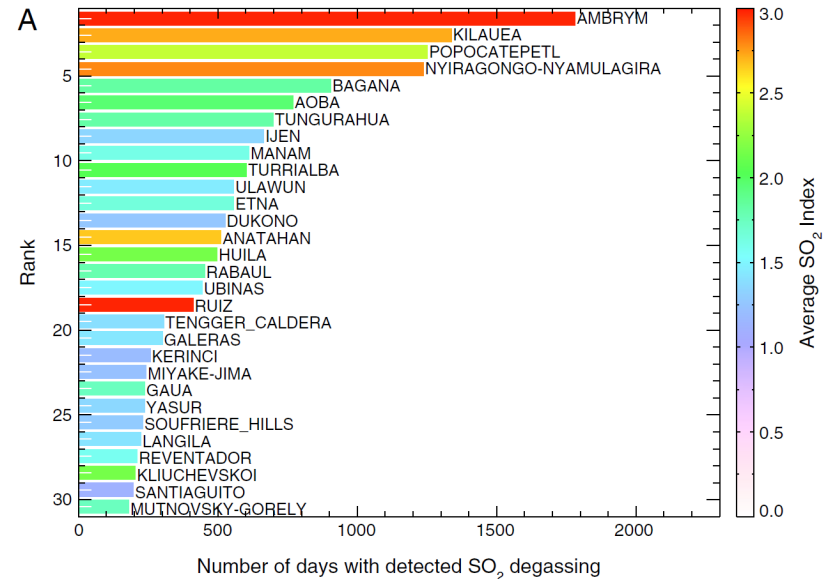


<sup>£</sup> NASA's SO<sub>2</sub> Emissions Group; \*Halmer et al. (2002), Bluth et al. (1993); Carn et al. (2016)

# Satellite-based SO<sub>2</sub> flux inventories in 2005-2015



Carn et al. 2017



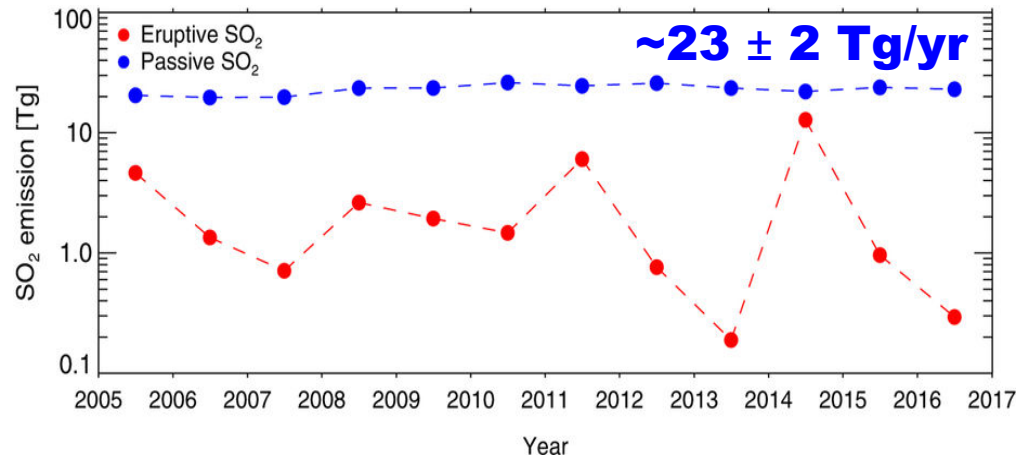
**Volcanic SO<sub>2</sub> emissions from 91 volcanoes in 2005–2015 derived from global satellite (OMI) measurements**

SCIENTIFIC REPORTS

OPEN

A decade of global volcanic SO<sub>2</sub> emissions measured from space

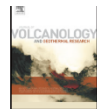
S. A. Carn<sup>1</sup>, V. E. Fioletov<sup>2</sup>, C. A. McLinden<sup>2</sup>, C. Li<sup>3,4</sup> & N. A. Krotkov<sup>4</sup>



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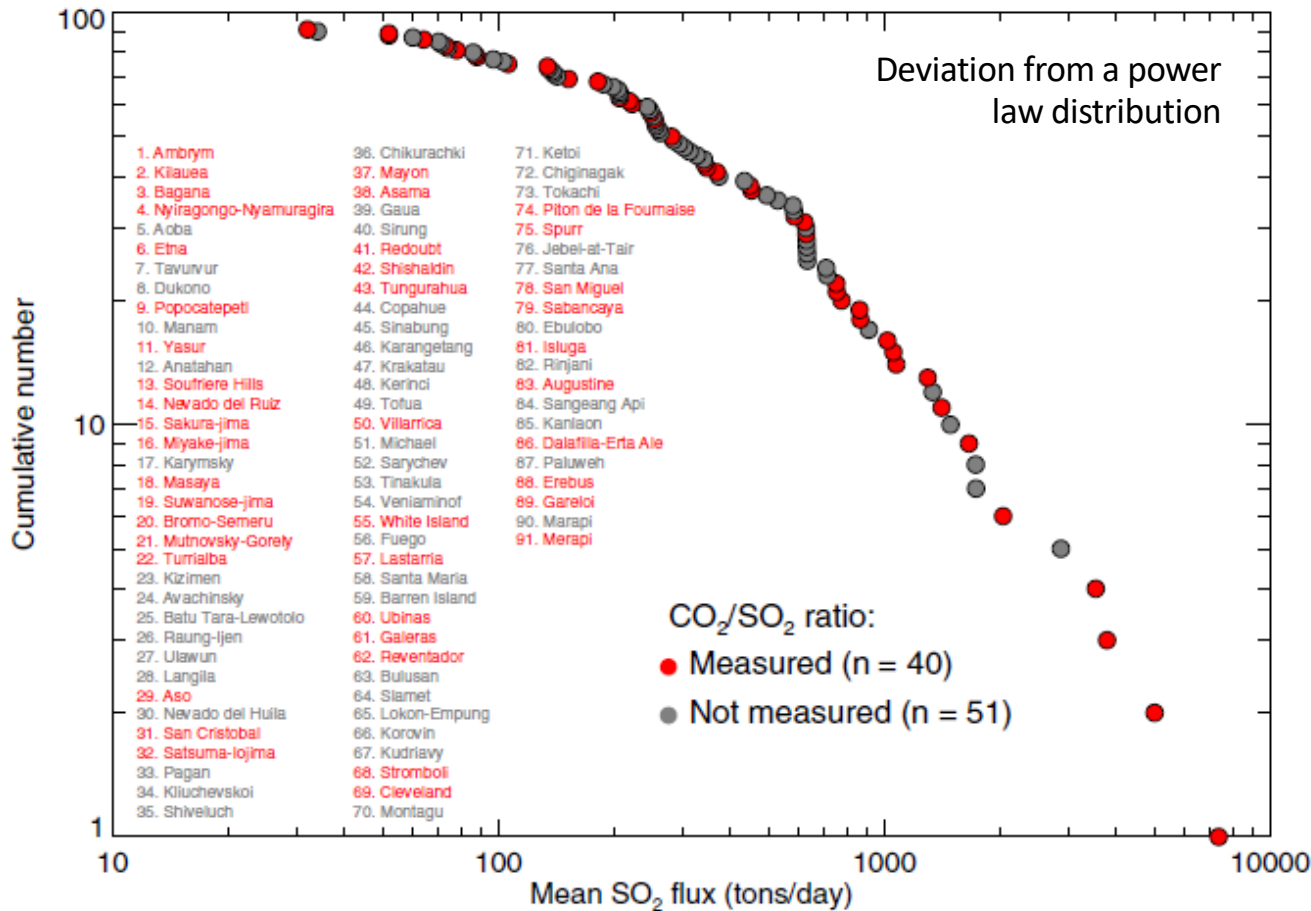
Review

Multi-decadal satellite measurements of global volcanic degassing

S.A. Carn<sup>a,b,\*</sup>, L. Clarisse<sup>c</sup>, A.J. Prata<sup>d</sup>



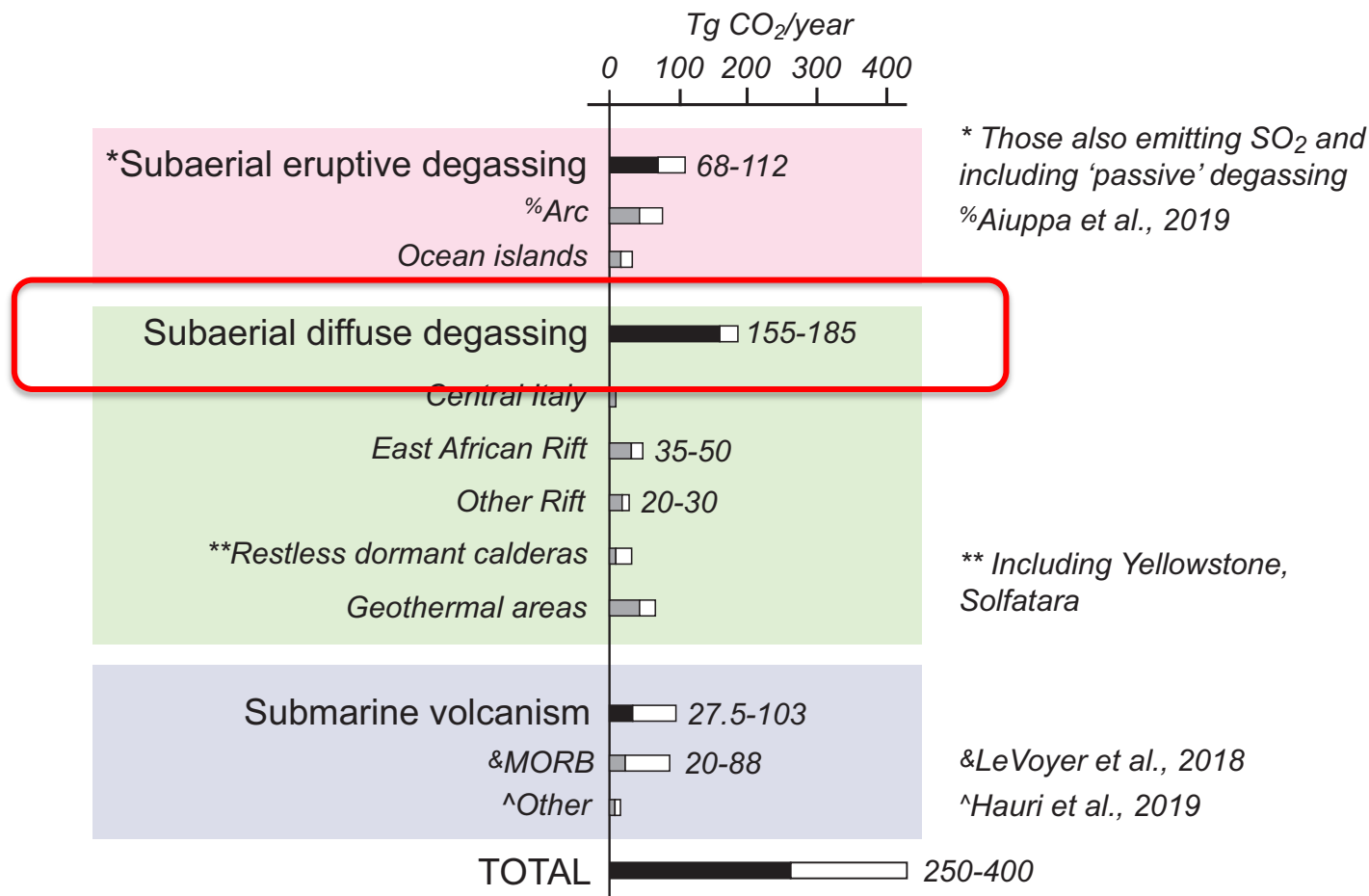
Strategy: a top few CO<sub>2</sub> and SO<sub>2</sub> volcanic outgassers dominate flux



10-12 volcanoes produce 95% of the sulfur flux to the atmosphere.

Carn et al., 2017

**Global Volcanic Carbon Budget**

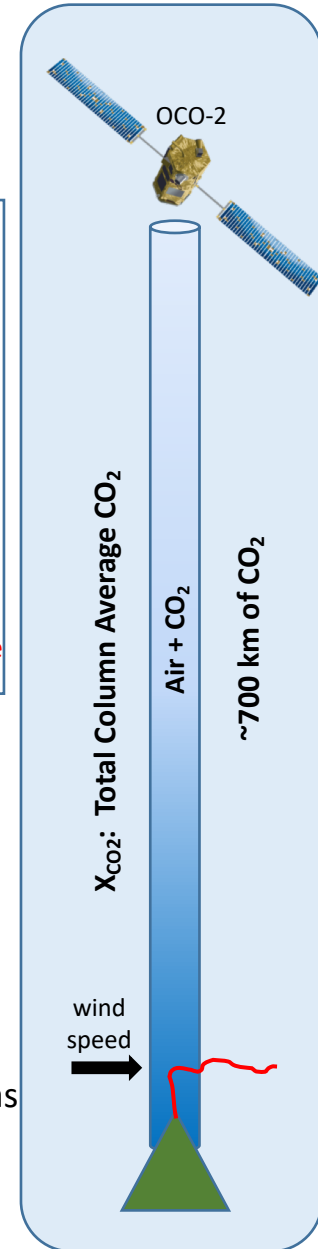
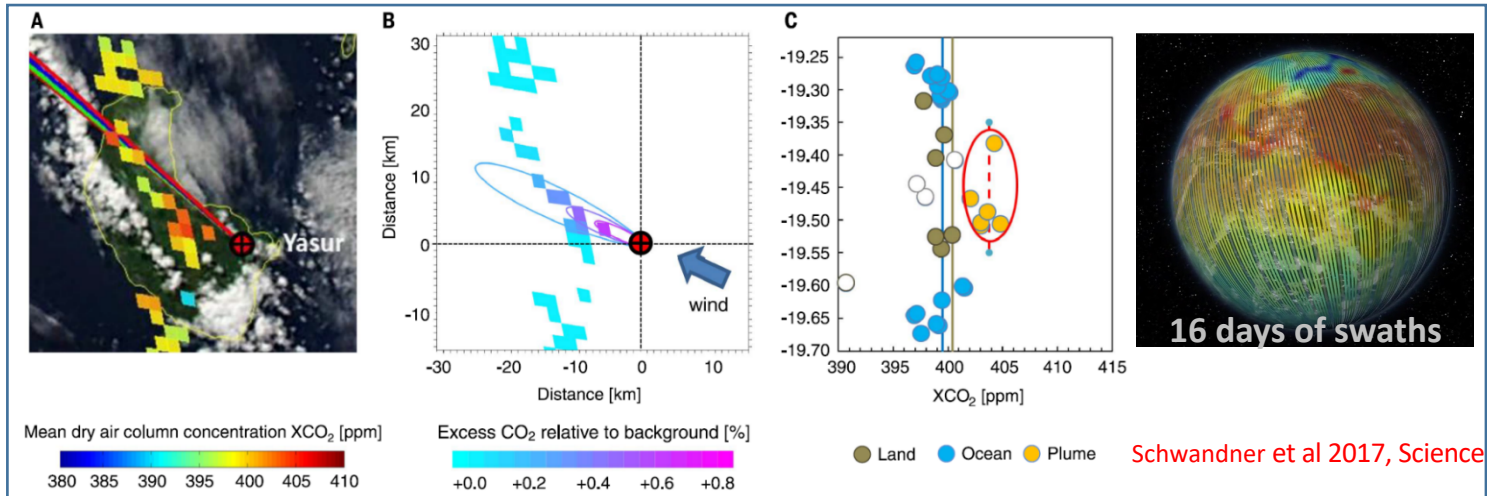


Large uncertainties

Werner et al., 2019, in press

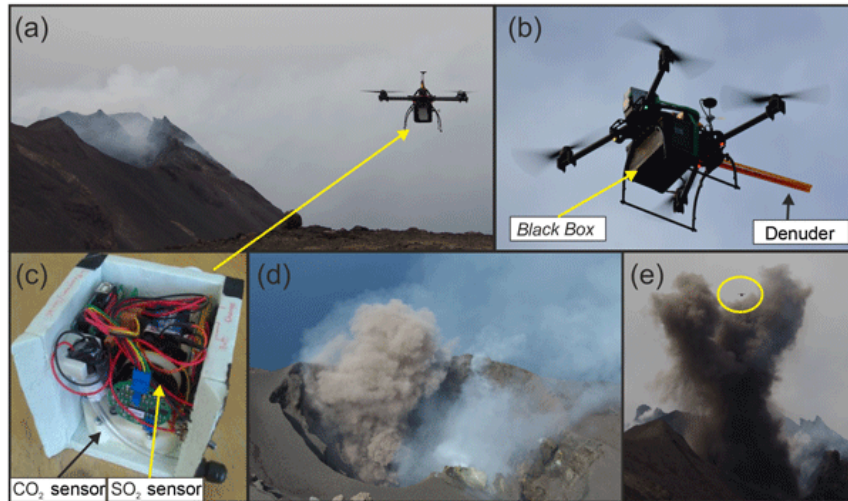
# The future: Direct detection/survey of volcanic CO<sub>2</sub> from space ?

- **Challenge:** “total column” average CO<sub>2</sub> ( $X_{CO_2}$ ) signal requires >99.75% precision (**1/410 ppm**)
- NASA and JAXA satellites (**OCO-2**, **GOSAT**) observed CO<sub>2</sub> over **>45 volcanoes since 2010**
- First space-borne (OCO-2) detection of volcanic CO<sub>2</sub> in 2015 at **Yasur**, Ambrym, Aoba (Vanuatu) (Schwandner et al., Science 2017)



- **Limitations:** OCO-2 detections are limited by sparse narrow-swath sampling pattern.
- New: **OCO-3** (launched to ISS May 4<sup>th</sup> 2019) will **cover several tens of volcanoes in dedicated mapping and target modes**, a significant improvement over OCO-2's capabilities.
- Problem with direct detections from space: Satellites can't **directly** detect 'mild' long-term precursory signals due to the dilution problem of  $X_{CO_2}$  measurements.
- **Solution:** Plants *glow and grow* with volcanic CO<sub>2</sub>: The **"Trees as sensors"** concept (Bogue, Schwandner, et al 2019, Biogeosciences). Plants grow and fluoresce more under mild CO<sub>2</sub> enhancements; measurable by remote sensing as increases in biomass (long-term response) and photosynthesis (short-term response).

# The future: using unmaned airborne platforms or drones for locally measuring volcanic degassing, even during eruptions (already ongoing)



**BlackSwift S2™ UAS**  
• Robust, simple to operate, scientific platform

**In Situ Atmospheric Probe**  
• Pressure, Temperature, Humidity  
• Three-Dimensional Winds

**Nephelometer**  
• Particle Size and Distribution

**Cameras**  
• EO/Thermal Images + Forward Video

**Dedicated Trace Gas Sensors**  
• SO<sub>2</sub>, CH<sub>4</sub>, H<sub>2</sub>S, CO<sub>2</sub>



**ABOVE**  
Aerial-based Observations of Volcanic Emissions

Today, a large (and young) international scientific community studying volcanic degassing. Gas monitoring now operated in many volcano Observatories!

## Commission on the Chemistry of Volcanic Gases (CCVG) - IAVCEI



- ❑ Inter-comparison of techniques and results in direct sampling of volcanic gases, widened to studies with ground and satellite remote sensing, in-situ automated devices, diffuse gas probes, airborne measurements, petro-chemical modelling, etc.
- ❑ A **Field Workshop every 3 years**, in different volcanic places, since 1982!
- ❑ Joint publications, a Newsletter (*Telegram from the Earth's Interior*)..

Kamchatka, 2011



Galapagos, 2017



Chile, 2014





Thank you for your attention

Ambrym volcano, Vanuatu arc  
Photo: P. Allard